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LEPISOSTEIFORM FISH (HOLOSTEI) GANOID SCALES FROM THE MIDDLE JURASSIC DEPOSITS OF UKRAINE

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Lepisosteiform Fish (Holostei) Ganoid Scales from the Middle Jurassic Deposits of Ukraine. Kovalchuk, O. M., Anfimova, G. V. — Gars (Lepisosteiformes) flourished in epicontinental seas throughout the world during the second half of Mesozoic and early Cenozoic. Cretaceous and Paleogene remains of these fishes are common in Europe while their Jurassic fossils are still relatively scarce. Here we re-describe in detail a series of ganoid scales from the latest Middle Jurassic (Callovian) deposits of Pekari (Cherkasy Region, Ukraine). These fossils were identified by Professor O. S. Rogovich in 1861 as those that belong to *Lepidotus mantellii* and *L. fittoni*. The referral of these scales to a certain species should be considered with caution because the described material lacks characters sufficient for identification it even to the genus or family. We consider to identify these fossils as *Lepisosteiformes* gen. et sp. indet. An overview of currently known Jurassic occurrences of lepisosteiform fishes is also presented in the paper. Key words: *Lepisosteiformes*, museum collection, morphology, Callovian, Europe.

Introduction

Order *Lepisosteiformes* is a small group of primitive neopterygian fishes comprising two families — *Lepisosteidae* with seven genera (only two of which — *Atractosteus* and *Lepisosteus* — remain extant), and †*Obaichthyidae* including two genera (Grande, 2010). In addition, seven extinct genera of lepisosteiform fishes (†*Arari-lepidotes*, †*Isanichthys*, †*Lepidotes*, †*Pliodetes*, †*Thaichthys*, and †*Scheenstia*) are indicated by incertae familiae (López-Arbarello, 2012). The current distribution of gars is restricted to North America, Central America and the Caribbean islands (Nelson et al., 2016). However, it was much wider during Mesozoic and early Cenozoic as evidenced by the presence of their fossils throughout the world (e. g. in South America, Africa, Europe and Asia; see Grande, 2010).

Lepisosteiform fishes appeared in the fossil record in the Early Jurassic and became diverse during the Late Jurassic and Early Cretaceous (López-Arbarello, 2012). However, little is known about the taxonomic composition and diversity of the Jurassic lepisosteiforms. Their remains in Europe are relatively scarce, but even if present, they are mostly unknown due to being “buried” in old natural history collections without a detailed description (Tretyak & Chervonenko, 2016).

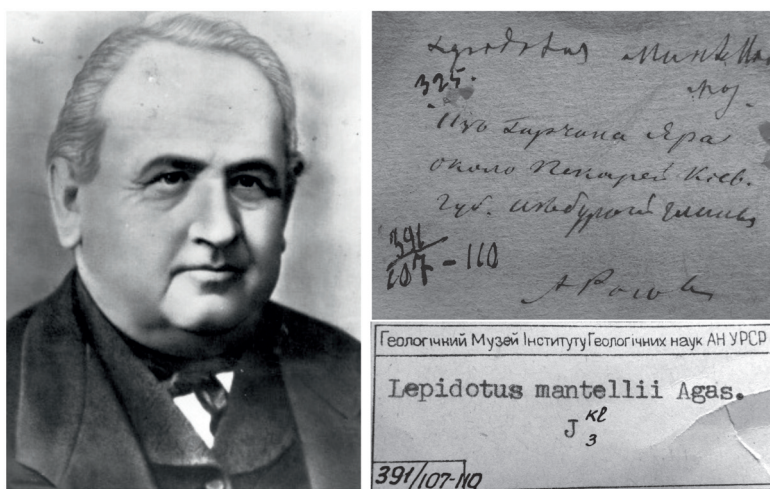


Fig. 1. Location of the Pekari fossil site on the map of Ukraine, the portrait of Professor O. S. Rogovich with the original (above) and additional labels (below) of the described material.

One of such finds is a series of ganoid scales briefly described by Rogovich (1861) from the Middle Jurassic deposits of Girchin quarry near Pekari, Cherkasy region, Ukraine (fig. 1). Rogovich (1861) assigned these scales to “*Lepidotus Mantellii*” and “*Lepidotus Fittoni*” based on their general similarity to scales of those species, which were previously described by Agassiz (1833–1844). The scales are “thin and thick, angular or completely rhomboidal in shape (depending on its position on the fish body)”, and their outer surface is “smooth or yields folds and dashes” (Rogovich, 1861: 67–68).

The aim of our study is to describe in detail the morphology of ganoid scales from Pekari, re-interpret this finding in terms of current stratigraphy and place it into a broad biogeographical context.

Geological setting, additional fauna and age estimation

The section from which the studied scales and numerous shark teeth were collected consists of “diluvium, Eocene green-grey sand and conglomerate, Cretaceous green sandstone, clayey-lignite intercalations, and Jurassic clay exposed at the base” (Rogovich, 1861: 7).

Middle Jurassic deposits in the Kaniv dislocation area are represented by a series of light-brown and dark-grey clay referring to the ammonite zone *Cadoceras* (*Paracadoceras*) *elatmae*, formerly defined as *Macrocephalites macrocephalus* zone (Kiselev & Ippolitov, 2011). A calcareous nannoplankton assemblage from these sediments was described by Matlaj (2016): *Biscutum dubium*, *Lotharingius contractus*, *L. crucicentralis*, *Stephanolithon speciosum*, *Watznaueria barnesae*, *W. britannica*, *W. fossacincta*, and *W. manivitiae*. These taxa represent the nannoplankton zone NJ12a/*Ansulasphaera helvetica* indicating the Early Callovian age (Matlaj,

2016). Based on this study, marine conditions within the Kaniv dislocation area already existed in the Early Callovian.

The ammonite assemblage of the studied region is represented by *Cadoceras* (*Paracadoceras*) *elatmae*, *Macrocephalites* cf. *verus*, *M. multicosatus*/*M. prosekensis*, *M. pavlowi*, *M. (Pleurocephalites)* cf. *terebratus*, and *Pseudocadoceras* (*Costacadoceras*) *mundum* (Kiselev & Ippolitov, 2011). This species list was supplemented by Gulyaev (2015) due to re-identification of some specimens from old museum collections and description of *Bullatimorphites* (*Kheiraceras*) *bullatus*. This taxon is common for the *Paracadoceras elatmae* biohorizon (Gulyaev, 2015), which is equal to the *Cadoceras suevicum* horizon in Germany (Mönnig, 2014). Gulyaev & Ippolitov (2013) reported about the findings of numerous Early Callovian belemnite remains from Kaniv district, namely of *Cylindroteuthis* s. str., “*Cylindroteuthis*” *kowalewi*, *Pachyteuthis* s. str., and *Communicobelus* represented by different morphs.

Material and methods

The examined series of scales is stored in the Department of Geology of the National Museum of Natural History, National Academy of Sciences of Ukraine, Kyiv, Ukraine. The scales were identified using diagnostic features based on comparisons with extinct and modern taxa (deposited in Virginia Institute of Marine Science, USA, Hungarian Natural History Museum, Budapest, Hungary, and Babeş-Bolyai University Cluj-Napoca, Romania) as well as on data from the literature (Agassiz, 1833–1844; Grande, 2010; López-Arbarello, 2012; Alvarado-Ortega et al., 2014; Sweetman et al., 2014; Pouech et al., 2015). The taxonomic hierarchy follows López-Arbarello (2012), and Nelson et al. (2016). Morphological description is presented here according to Grande (2010), with reference to other publications (e.g., Kerr, 1952; Thompson & McCune, 1984; Kumar et al., 2005; Alvarado-Ortega et al., 2014; Garbelli & Tintori, 2015; Kyselevych & Kovalchuk, 2019). The scales were measured by an electronic caliper, and photographed using a Leica M168C camera in Schmalhausen Institute of Zoology, National Academy of Sciences of Ukraine.

Systematic paleontology

Class ACTINOPTERYGII Cope, 1887 sensu Rosen et al., 1981

Subclass NEOPTERYGII Regan, 1923 sensu Nelson et al., 2016

Infraclass HOLOSTEI Müller, 1844 sensu Grande, 2010

Subdivision GINGLYMODI Cope, 1872 sensu Grande, 2010

Order LEPISOSTEIFORMES Hay, 1929 sensu López-Arbarello, 2012

Lepisosteiformes gen. et sp. indet.

Lepidotus mantellii Ag.: Rogovich, 1861, pp. 67–68; table IX, figs 27–34.

Lepidotus fittoni Ag.: Rogovich, 1861, p. 68; table IX, figs 35–39.

Material. Seven ganoid scales, Nos. 391/107a-d, 391/108-10.

Locality and age. Pekari (49°42' N, 31°33' E), Cherkasy Region, Ukraine; Middle Jurassic, Early Callovian (J₂kl₁), ca. 166–165 Ma.

Description. Ganoid scales are represented by three morphotypes.

Morphotype 1 (specimen No. 391/109; fig. 2, A₁-A₂). The scale plate is elongated, smooth and narrow, irregularly rectangular in shape; its total length is 15.0 mm, width — 8.0 mm. The thin ganoine layer covers the entire outer surface of the scale. Edges of the scale are smooth and rounded (posterodorsal angle is 103°, while the posteroventral angle equals 84°). The scale is the thickest along the ridge which is parallel to both its anterior and posterior edges. There is a weak serration (four small denticles with blunt tips) forming a shallow ridge on the dorsal edge, between the dorsal peg (*dp*) and posterior margin of the scale. The anterodorsally oriented dorsal anterior process (*dap*) is strong, stout and sharp while the ventral anterior process (*vap*) is totally reduced. The space between the anterior processes is filled with a thin plate. The dorsal peg is shifted toward the *dap* base. Posterior edge of the scale is narrow; the ventral edge forms a long shallow ridge.

Morphotype 2 (specimens No. 391/107a, fig. 2, B₁-B₂; No. 391/107b, fig. 2, C₁-C₂; No. 391/110, fig. 2, D₁-D₂). The scales are narrow and elongated, their length varies in the range of 12.0–13.0 mm, width of the scale plate — 4.0–5.0 mm. There is a very strong dorsal

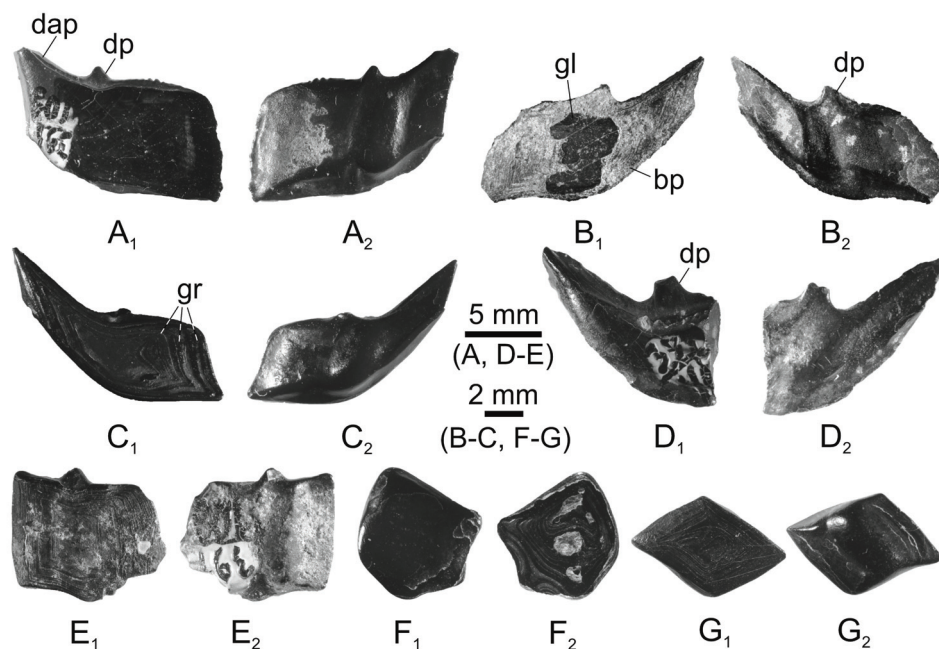


Fig. 2. Ganoid scales of *Lepisosteiformes* gen. et sp. indet. from Pekari: A₁-A₂ — No. 391/109; B₁-B₂ — No. 391/107a; C₁-C₂ — No. 391/107b; D₁-D₂ — No. 391/110; E₁-E₂ — No. 391/108; F₁-F₂ — No. 391/107d; G₁-G₂ — No. 391/107c. Outer view in A₁, B₁, C₁, D₁, E₁, F₁, G₁, inner view in A₂, B₂, C₂, D₂, E₂, F₂, G₂. Abbreviations: *bp*, basal plate; *dap*, dorsal anterior process; *dp*, dorsal peg; *gl*, ganoin layer; *gr*, ganoin ridges.

anterior process (5.0–8.0 mm in length) with a widened base and pointed tip. The ventral anterior process is absent as well, and the dorsal peg is comparatively reduced. The latter is located just near the *dap* base. All the scales are covered with a thin ganoin layer, however the specimen No. 391/107a (fig. 2, B₁-B₂) partially lacks this cover (due to postmortem redeposition), and therefore its basal plate is uncovered. The scales are the thickest on the vertical of the *dp*. There are two shallow ridges along the margins of the *dap*, as well as one longitudinal ridge on its outer surface. The posterodorsal edge of all scales is rounded. There is a concave groove starting from the proximal third of the *dap* up to the posterior half of the *dp* in specimen No. 391/110 (fig. 2, D₁-D₂).

Morphotype 3 (specimen No. 391/107c, fig. 2, G₁-G₂). The small scale (total length 8.5 mm, width 6.0 mm) is clearly rhomboidal in shape. Its outer surface is smooth and slightly convex. As other scales in the described series, this specimen is narrower toward the non-serrated edges being the thickest in its central part. The anterior processes are not expressed, and the *dp* is absent.

Two other scales (No. 391/108, fig. 2, E₁-E₂; No. 391/107c, fig. 2, F₁-F₂) are represented by small rounded and slightly concave fragments with concentric rings on the smooth outer surface, weakly expressed *dap* and the *dp*.

C o m p a r i s o n . The ganoid scales from Pekari resemble those in representatives of the order *Lepisosteiformes* in overall morphology, including the presence of two processes forming a rostro-caudal or longitudinal articulation (Cavin et al., 2009) and the reduction of the ventral anterior process diagnostic of the suborder *Lepisosteioidei* (López-Arbarello, 2012). However, these characters differ from those in the lepisosteiform crown taxa, whose scales have the peg-and-socket vertical articulation reduced or absent in the superfamily *Lepisosteioidea* and a strong posteriorly directed spine in the suborder *Lepisosteioidei* (López-Arbarello, 2012; Alvarado-Ortega et al., 2014).

The general shape of some scales from Pekari (morphotypes 2 and 3) is similar to the scales of ?*Scheenstia* sp. described by Sweetman et al. (2014: fig. 13, E-F, p. 891) from the

Lower Cretaceous deposits of England. Both of them are equal in size and share the same type of articulation. However, the specimens Nos. 391/109, 391/107a, and 391/107b are more elongated and have a stronger dorsal anterior process. Besides, the dorsal edge of the scales in ?*Scheenstia* sp. is non-serrated, and the scale plate seems to be deeper.

A scale fragment of *Scheenstia mantelli* from the Late Jurassic of Cherves-de-Cognac in France (Pouech et al., 2015: Fig. 3O) resembles No. 391/107d (fig. 2, F₁-F₂). Other known scale specimens of *Scheenstia* (Murray, 2000: Fig. 6B, p. 127; Pouech et al., 2015: Fig. 3R, p. 36; Haddoumi et al., 2016: Fig. 10A, p. 298; Kyselevych & Kovalchuk, 2019: fig. 2, etc.) share the same morphology and similar ganoin ridges at the outer surface, but have two well-developed anterior processes as well as shorter and deeper scale plate. We can say the same about the scales of *Lepidotus* figured by Agassiz (1833–1844) and described in detail by Woodward (1919). It should be noted that the morphological differences between *Lepidotus* and *Scheenstia* are not completely resolved because not all the known *Lepidotus* species were considered in the study of López-Arbarello (2012).

The scales of lepisosteiform fishes show significant differences along the trunk of the same individual (see Grande, 2010; López-Arbarello, 2012; Alvarado-Ortega et al., 2014). The attribution of isolated ganoid scales with a certain species (as it was presented by Rogovich, 1861) should be considered with caution because such material often lacks sufficient diagnostic characters. In our case, it is not possible to make an accurate identification and decide whether these scales represent one or two species. We decide to assign the specimens from Pekari as *Lepisosteiformes* gen. et sp. indet.

Discussion

The oldest reliable lepisosteiform remains are known from the Lower Toarcian deposits (182 Ma) of Grimmen in Germany (Böhme & Ilg, 2003). The finding of *Lepisosteiformes* gen. et sp. indet. in Pekari is the sole Callovian record of this group in Eastern Europe and one of its few occurrences currently known for this age. Another one is *Isanichthys latifrons* (Woodward, 1893) from the Middle Callovian of Oxford Clay Pit in England (Woodward, 1893), as well as the scales identified as *Scheenstia* sp. from the Middle Jurassic of Guelb el Ahmar, Morocco (Haddoumi et al., 2016). There is a short report of Callovian fish remains found in Sarykamyskai 1 (Kyrgyzstan) and assigned to *Lepisosteus?* sp. (Böhme & Ilg, 2003), however this record is doubtful (Grande, 2010), and *Lepisosteiformes* are regarded to be absent in the general faunal list from this locality (Averianov et al., 2005).

As for other Jurassic (but non-Callovian) lepisosteiform occurrences, most of them are known from Europe (Böhme & Ilg, 2003). Those are remains of *Scheenstia laevis* (Agassiz, 1833–1844) from the Kimmeridgian/Tithonian of Cerin, Ain and *Scheenstia maximus* (Wagner, 1863) from Lot-et-Garonne, both in France (Agassiz, 1833–1844; Sauvage, 1902). Swiss record of Jurassic lepisosteiform fossils is represented by teeth and scales of *Scheenstia* sp. from the Kimmeridgian of Bois de Sylleux and Tchâfoué (López-Arbarello & Sferco, 2011; López-Arbarello, 2012).

Lepisosteiform fossils, mostly represented by the genus *Scheenstia* López-Arbarello & Sferco, 2011, are quite common in lithographic limestones of the Solnhofen and Mörnsheim formations (Late Jurassic) in Germany. *Scheenstia maximus* (Wagner, 1863) was reported from Eichstätt, Kellheim (Wagner, 1863), and Langenthaltheimer Haardt (Böhme & Ilg, 2003). López-Arbarello & Sferco (2011) described *Scheenstia zappi* from the Kimmeridgian of Schamhaupten. *Scheenstia decoratus* (Wagner, 1863) is known from the Lower Tithonian beds of Solnhofen (López-Arbarello, 2012). Other species of the genus *Scheenstia* (*S. degenhardti* (Branco, 1885), *S. hauchecornei* (Branco, 1887)) were reported from the Upper Berriasian lagoon deposits of Obernkirchen near Bückeberg (Böhme & Ilg, 2003).

There are a few Jurassic localities yielding the lepisosteiform remains outside of Europe: Qijiang in China (Tithonian/Kimmeridgian: *Beiduyu quijiangensis* Murray et al., 2015), Phu Nam Jun in Thailand (Tithonian: *Isanichthys palustris* Cavin & Suteethorn, 2006; *Thaichthys buddhabutrensis* (Cavin et al., 2003)), and Yosobé near Tlaxiaco in Mexico (Kimmeridgian: *Nhanulepisosteus mexicanus* Brito, Alvarado-Ortega & Meunier, 2017).

Every new find of lepisosteiform fossils (even described in open nomenclature) is important for precise reconstruction of the biogeographical history of this group. The revision of old natural history collections is of great interest because it allows re-discovering some valuable but unknown specimens, and describing them in more detail using the newest methodology.

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