

## MIOCENE QUANTITATIVE CALCAREOUS NANNOFOSSIL BIOSTRATIGRAPHY FROM SOUTHERN APENNINES FOREDEEP DEPOSITS AND MEDITERRANEAN DSDP SITE 372

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**Riassunto.** Sui campioni raccolti in nove successioni mioceniche rappresentate prevalentemente da depositi calciclastici e silicoclastici, riferibili alle unità irpine esterne ed affioranti al margine esterno dell'Appennino meridionale (Italia meridionale), è stato condotto uno studio biostratigrafico quantitativo basato sulle associazioni a Nannofossili calcarei. Le analisi quantitative hanno permesso di riconoscere, anche in queste successioni, gran parte degli eventi biozonali noti per l'area mediterranea in accordo con gli schemi biostratigrafici di Fornaciari & Rio (1996) e Fornaciari et al. (1996) per il Miocene inferiore e medio, e di Theodoridis (1984) per il Miocene superiore.

Alcuni tra gli eventi biozonali utilizzati in letteratura, come ad esempio la "first common occurrence" di *Helicosphaera walbersdorfensis* o la "last common occurrence" di *Calcidiscus premacintyreii*, sono stati però di difficile definizione; nell'intervallo MNN6b/7 la "first common occurrence" di *Calcidiscus macintyreii* rappresenta invece, nelle successioni analizzate, un evento più significativo rispetto alla "last common occurrence" di *Calcidiscus premacintyreii*.

Sono stati inoltre identificati nuovi eventi biostratigrafici che migliorano la risoluzione degli schemi biozonali utilizzati: nelle Zone MNN4b e MNN5a è stato riconosciuto un significativo intervallo di paracme di "small" *Reticulofenestra pseudoumbilicus*; il "paracme beginning" ed il "paracme end" di questa specie sono eventi utili nella correlazione delle successioni studiate. Inoltre, la presenza di *Reticulofenestra pseudoumbilicus* è stata riconosciuta a partire dalla Zona MNN2b, in un intervallo quindi molto inferiore rispetto alla nota "first common occurrence" della specie, che definisce la base della Zona MNN6b.

Le distribuzioni osservate nelle successioni dell'Appennino meridionale sono state inoltre confrontate con i risultati ottenuti da uno studio quantitativo svolto sul Pozzo DSDP 372 ubicato nel Mediterraneo occidentale: ciò ha permesso di confermare il valore biostratigrafico di tali distribuzioni nell'area mediterranea.

**Abstract.** A quantitative biostratigraphic study based on calcareous nannofossil assemblages was carried out in nine Miocene calcareous and siliciclastic foredeep sections, cropping out in the outer part of the Southern Apennines and generally ascribed to the external Irpinian units. Several biozonal events were recognised by means of quantitative analyses, according to the biostratigraphic schemes of Fornaciari & Rio (1996) and Fornaciari et al. (1996) for the lower and middle Miocene and of Theodoridis (1984) for the upper Miocene.

In the lower and middle Miocene interval some biohorizons such as the first common occurrence of *Helicosphaera walbersdorfensis* and last common occurrence of *Calcidiscus premacintyreii* are not always detectable in the studied sections and the first common occurrence of *Calcidiscus macintyreii* appears to be a better biohorizon in the MNN6b/7 than the last common occurrence of *C. premacintyreii*.

New data on abundance patterns of selected calcareous nannofossils have been collected, thus improving the biostratigraphic resolution of the zonal schemes: a paracme interval of small *Reticulofenestra pseudoumbilicus* was noted within Zone MNN4b and MNN5a; the beginning and the end of the paracme are useful events in the studied sections for stratigraphic correlations. Moreover the occurrence of *R. pseudoumbilicus* has been recorded from Zone MNN2b on upwards, fairly below the FCO of the species that defines the base of Zone MNN6b.

The distributions observed in the on-land sections are consistent with those recognised in DSDP Site 372 located in the western Mediterranean Sea, confirming their potential biostratigraphic utility within the Mediterranean region.

### Introduction.

Recent studies have been focused on the improvement of the Miocene calcareous nannofossil biostratigraphy within the Mediterranean region (Theodoridis, 1984; Negri, 1989; Di Stefano, 1993; Fornaciari & Rio, 1996; Fornaciari et al., 1996). In particular, Fornaciari & Rio (1996) and Fornaciari et al. (1996) proposed new biostratigraphic schemes based on abundance fluctuations of selected calcareous nannofossil species and these schemes provide a higher biostratigraphic resolution in comparison with the standard zonations of Martini (1971) and Okada & Bukry (1980). In order to provide new data on the applicability of the recently proposed schemes in different areas, a biostratigraphic study based on calcareous nannofossils in Miocene foredeep deposits has been performed. Such deposits are widespread in the Southern Apennines and their biostratigraphic dating is also very useful for the understanding and modelling of

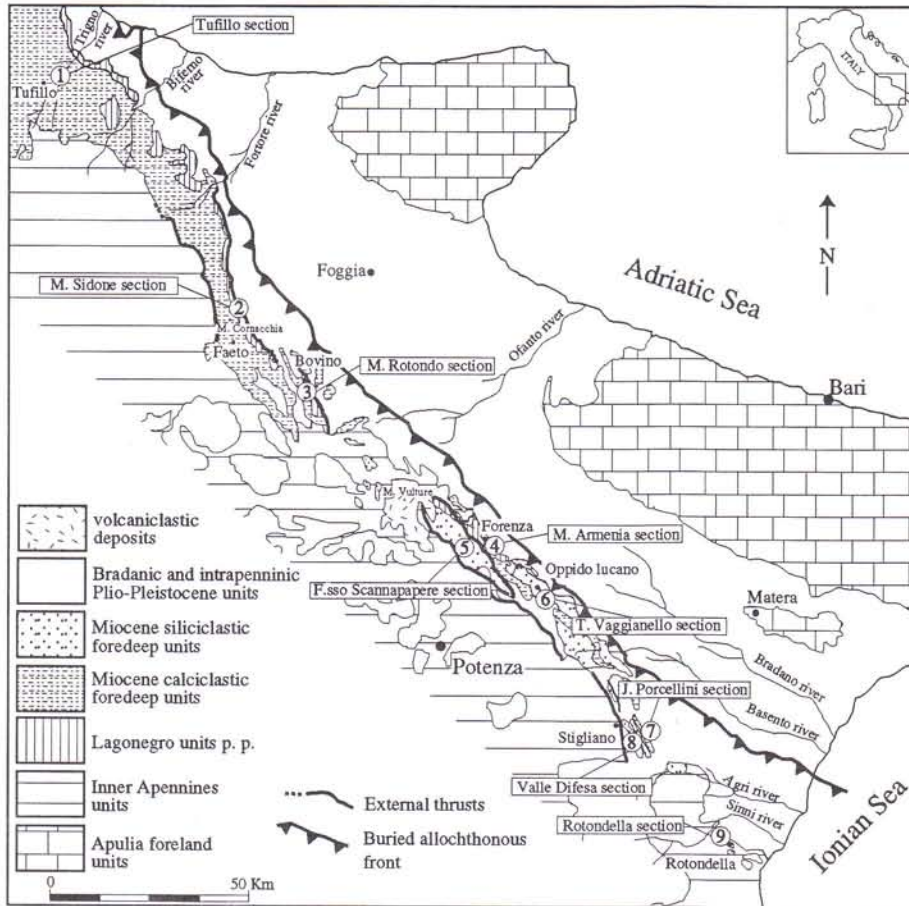


Fig. 1 - Location of the studied sections in Southern Apennines; geological map from Bonardi et al. (1988) and Bigi et al. (1989), modified.

ascribed to the outer Irpinian units (Pescatore, 1978; Dazzaro & Rapisardi, 1983; Pescatore, 1988).

The herein discussed calcareous sections (Tufillo, Monte Sidone, Monte Rotondo, Monte Armenia and Rotondella) are mostly referred to the Flysch di Faeto Formation (Crostella & Vezzani, 1964) and to the Tufillo Formation (Selli, 1962) and crop out mostly between the Trigno river and the Monte Vulture. The siliciclastic sections (Fosso Scannapapere, Torrente Vaggiannello, Jazzo Porcellini, Valle Difesa) are ascribed to the Serra Palazzo Formation (Selli, 1962) and are widespread in the southern part of the studied area, between the Monte Vulture and Rotondella.

several geological problems (Gallicchio, 1996; Gallicchio & Maiorano, 1997; Gallicchio, 1997).

The biostratigraphic data on the calcareous nanofossil assemblages have been collected by means of quantitative analyses; these are of great utility for the recognition of significant abundance patterns and provided accurate age determinations as well as detailed stratigraphic correlations among the studied sections (Maiorano, 1996).

The investigation was extended to DSDP Site 372, located in the western Mediterranean Sea and representing a reference deep-sea section for the Mediterranean middle Miocene biostratigraphy (Müller, 1978; Theodoridis, 1984; Negri, 1989; Fornaciari et al., 1996). The re-examination of DSDP Site 372 in this study was performed in order to attempt a comparison with the quantitative results obtained from the Southern Apennines sections and to provide useful information on the Miocene Mediterranean calcareous nanofossil biostratigraphy.

#### Geological framework.

The investigated Miocene carbonatic and siliciclastic turbidite sections crop out in the outer part of the Southern Apennines, between the Trigno river and the neighbourhood of Rotondella (Fig. 1); they are generally

These units were deposited in a NW-SE foredeep domain whose western margin consists of deformed Apennine thrust belt, whereas the eastern one is represented by the undeformed Apulia Foreland (Fig. 1). The studied sections lie stratigraphically on the Oligocene-Miocene portion of the Lagonegro units (*sensu* Cocco et al., 1974) represented by the Numidian Flysch Formation (Ogniben, 1963) or by the Calcareni, marne ed argille di M. Sidone Formation (Senatore, 1988). They are overlaid by the Marne argillose del Toppo Capuana Formation (Crostella & Vezzani, 1964).

Further lithostratigraphic and sedimentologic details on the Southern Apennines Miocene foredeep deposits are reported in many papers to which the reader is referred (Ippolito & Lucini, 1957; Palmentola, 1969, 1970; Palmentola et al., 1967, Pieri & Walsh, 1967; Bonenzi et al., 1968; Ogniben, 1969; Centamore, 1969; Centamore et al., 1971; Rapisardi & Walsh, 1978; Pescatore, 1978, 1988; Ciaranfi et al., 1980; Mostardini & Merlini, 1986; Di Nocera & Torre, 1987; Bonardi et al., 1988; D'Argenio, 1988; Dazzaro et al., 1988; Sgrosso, 1988; Russo & Senatore, 1989; Patacca et al., 1990, 1992; Loiacono & Sbarra, 1991; Boiano et al., 1994; Sbarra, 1995; Gallicchio, 1996). In particular a detailed lithostratigraphic and structural framework of the studied sections is documented in Gallicchio (1996).

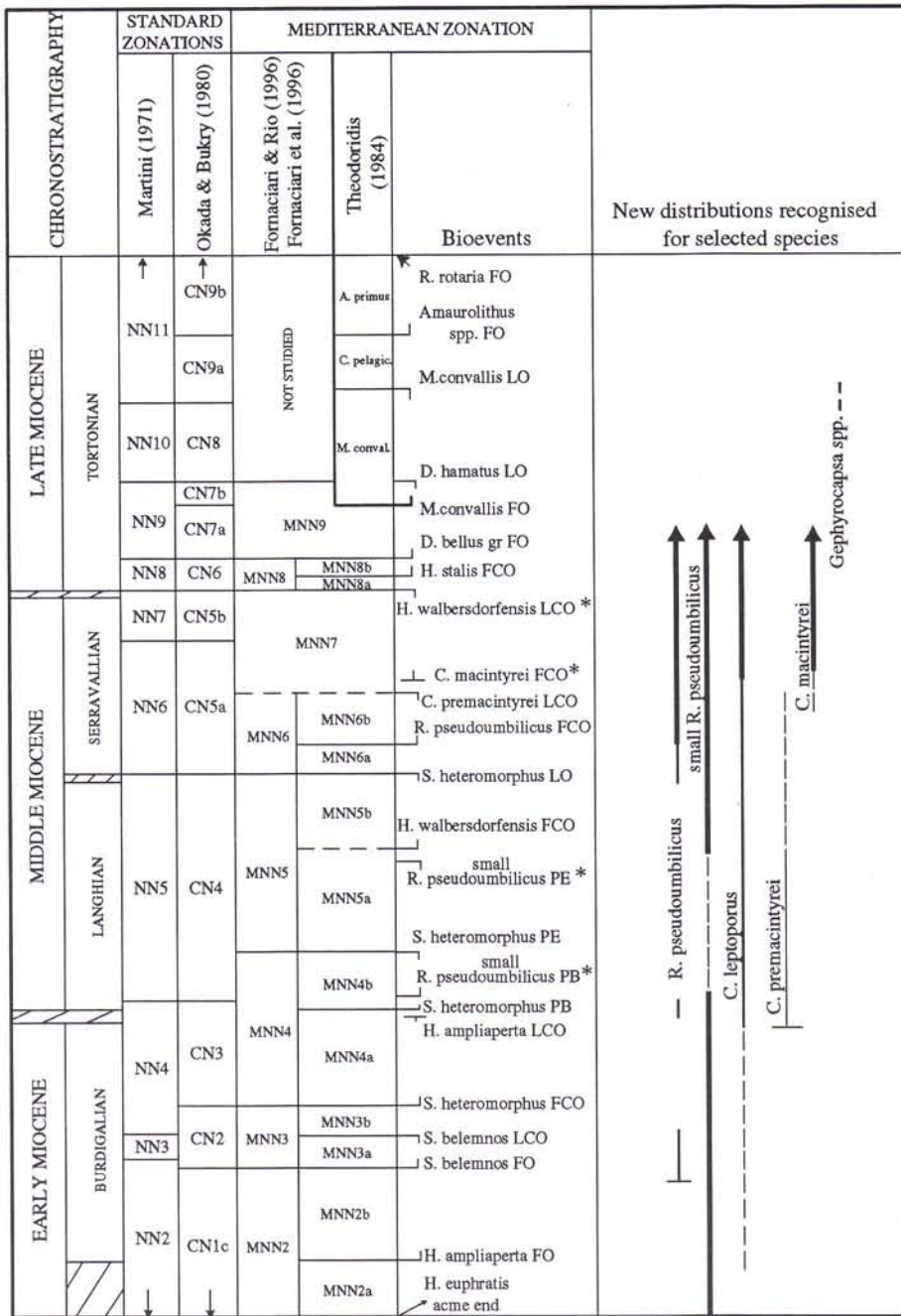


Fig. 2 - Biostratigraphic scheme and distribution of some marker species proposed for the Mediterranean region and comparison with standard zonations. Undetectable biozonal boundaries are traced with dashed lines; \* = new biostratigraphic events proposed in the present work. Chronostratigraphy is drawn according to Fornaciari and Rio (1996) and Fornaciari et al. (1996); buffer intervals indicate that at present no formal definition of the boundary stratotype sections is available.

in all the samples in order to recognise the presence of rare species. Moreover, in the upper Miocene interval, a number of about 10,000 specimens has been scanned in order to get significant abundance fluctuations of *Discoaster* spp. and *Amaurolithus* spp. and the abundance patterns have been plotted as number of specimens/mm<sup>2</sup>. The diagrams of the quantitative distribution of the most significant species are shown for each section; distribution charts of the total assemblages are included.

Calcareous nannofossils in the studied material are often poorly preserved. The total abundance in the samples is generally no more than 7-10 specimens per field of view and it is often represented by 3-5 specimens only. Reworked Cretaceous and Cenozoic nannofossils are present. In the siliciclastic sections the number of reworked specimens on 500 nannofossils is generally greater than in the calcareous ones and may reach value up to 40% of the assemblage. However the reworked taxa did not prevent the recognition of the biozonal events; quantitative patterns are in fact comparable with those observed for most of the species in pelagic or hemipelagic sequences.

**Materials and Methods.**

Samples for calcareous nannofossil analyses were prepared from a suspension of unprocessed sediment and water. The solution was spread on a cover glass and dried, then mounted on a glass slide. Light microscope techniques were used at 1000X magnification.

Quantitative analyses were performed according to Rio et al. (1990b), in order to detect the abundance patterns of the marker species. These methods have been already tested in turbidite sediments (Fornaciari & Labaume, 1992; Fornaciari & Rio, 1996; Fornaciari et al., 1996).

Most of the quantitative data were collected by counting 500 specimens in the nannofossil population greater than 4 microns. A number of 30-100 taxonomically related specimens have also been counted in order to document the abundance fluctuations of helicoliths. Abundance patterns of *Discoaster variabilis-exilis* and *Discoaster deflandrei* were plotted as number of specimens/mm<sup>2</sup> and are relative to about 1500 specimens of the total nannofossil assemblage. A supplementary qualitative analysis on about 3000 specimens was performed

The following biometrical definitions have been adopted in this study:

- small *R. pseudoumbilicus*: reticulofenestrids 5-7 micron in size;
- R. pseudoumbilicus*: reticulofenestrids >7 microns (Raffi & Rio, 1979);
- large *R. pseudoumbilicus* reticulofenestrids >12 microns;
- Calcidiscus macintyreii*: circular *Calcidiscus* ≥11 microns (Rio et al., 1990a);
- Coccolithus miopelagicus* >13 microns (Perch-Nielsen, 1985).

**Adopted zonal scheme.**

In the past years various authors (Müller, 1978; Ellis, 1979; Ellis & Lohman, 1979) pointed out the limits of the standard biostratigraphic schemes of Martini (1971) and of Bukry (1973, 1975) within the Mediterranean region, due to the absence or the rarity of several

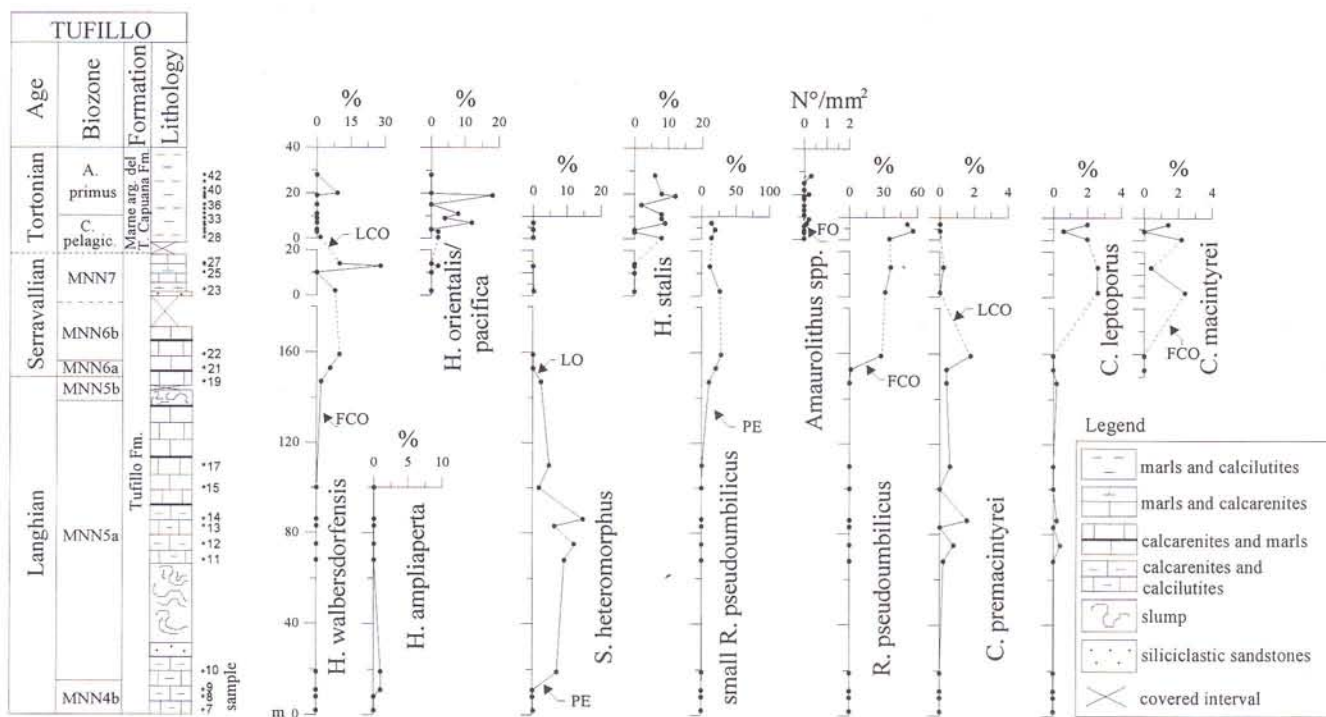


Fig. 3 - Abundance patterns of selected calcareous nannofossils at the Tuffillo section. FO = First Occurrence; LO = Last Occurrence; FCO = First Common Occurrence; LCO = Last Common Occurrence; PB = Paracme Beginning; PE = Paracme End.

marker species such as *Discoaster druggii*, *Triquetrorhabdulus carinatus*, *Catinaster coalitus* and *Discoaster kugleri*. Some alternative biozonal schemes have been proposed for the Mediterranean region (Cati & Borsetti, 1970; Ellis, 1979; Ellis & Lohman, 1979; Theodoridis, 1984; Negri 1989; Di Stefano, 1993; Fornaciari & Rio, 1996; Fornaciari et al., 1996) in order to improve the low resolution of the standard schemes for regional correlations.

In this study the most recent biostratigraphic schemes (Fornaciari & Rio, 1996; Fornaciari et al., 1996) proposed for the lower and middle Miocene Mediterranean sediments are followed. They are based on quantitative analyses on calcareous nannofossil assemblages carried out on various Italian sections from different depositional settings, as well as on the Langhian and Serravallian stratotype sections. Some of the events are from the zonal schemes of Martini (1971) and Bukry (1973, 1975) and of Theodoridis (1984). The new proposed biohorizons are based not only on FO (first occurrence) and LO (last occurrence), but on significant abundance fluctuations of some species, which allow to recognise supplementary events such as FCO (first common occurrence), LCO (last common occurrence), PB (paracme beginning) and PE (paracme end).

In Fig. 2 a comparison between the standard zonations and the Mediterranean biostratigraphic scheme adopted in this work is shown. In the zonal schemes of Fornaciari & Rio (1996) and Fornaciari et al. (1996) a dashed line was used to indicate those biohorizons not always recognised in the studied material, such as the

FCO of *H. walbersdorfensis* and the LCO of *C. premacintyreii*. Additional events as well as significant distributions of species recognised in this study have been included.

For the upper Miocene interval the biozonal scheme of Theodoridis (1984) was followed. The biozones of Theodoridis (1984) provide a better biostratigraphic resolution compared to the standard schemes and to what is known from the Tortonian-Messinian stratotype sections (Martini, 1975; Rio et al., 1976; Mazzei, 1977). Recent studies on the Falconara section (Sprovieri et al., 1996b) confirmed the applicability of the scheme of Theodoridis in the Tortonian-Messinian interval.

Although at present there is no general agreement on the formal definition of the Miocene boundary stratotype sections, the proposal discussed in Fornaciari et al. (1996) and Fornaciari et al. (1997) has been followed in this study in the lower and middle Miocene interval. The FO of *Reticulofenestra rotaria* has been considered for the recognition of the Tortonian/Messinian boundary, according to Langereis et al. (1984).

### Biostratigraphic Results.

#### Tuffillo section.

This section was sampled mostly along the Trigno river (154 IV SE, topographic map of Italy) and is referable to the Tuffillo Formation. It mainly consists of calcarenites, calcilutites and marls which lie stratigraphical-

Burdig.		Langhian		Serravallian		Tortonian		Age	
MNIN 3b		MNIN 5a		MNIN 7		A. primus *		Fornaciari & Rio (1996) Fornaciari et al. (1996) *Theodoridis (1984)	
Numidian Flysch Fm.		Tufillo Fm.		Tufillo Fm.		Marne argillose del T. Capuana Fm.		Formation	
1		7		18		30		Sample	
2		8		19		31		Total abundance	
3		9		20		32		Preservation	
4		10		21		33		Amaurolithus primus	
5		11		22		34		Amaurolithus spp.	
6		12		23		35		Calcidiscus leptoporus	
7		13		24		36		C. macintyreii	
8		14		25		37		C. premacintyreii	
9		15		26		38		Cryptococcolithus mediaperforatus	
10		16		27		39		Coccolithus miopelagicus	
11		17		28		40		C. pelagicus	
12		18		29		41		Coronocyclus nitescens	
13		19		30		42		Cyclicargolithus floridanus	
14		20		31		43		Discoaster berggrenii	
15		21		32		44		D. berggrenii/quinquerramus	
16		22		33		45		D. brouweri	
17		23		34		46		D. deflandrei	
18		24		35		47		D. intercalaris	
19		25		36		48		D. prepentaradiatus	
20		26		37		49		D. surculus	
21		27		38		50		D. triradiatus	
22		28		39		51		D. variabilis-exilis	
23		29		40		52		Discoaster spp.	
24		30		41		53		Geminilithella rotula	
25		31		42		54		Helicospaera ampliaptera	
26		32		43		55		H. carteri	
27		33		44		56		H. euphratis	
28		34		45		57		H. orientalis/pacifica	
29		35		46		58		H. scissura	
30		36		47		59		H. stalis	
31		37		48		60		H. walbersdorfensis	
32		38		49		61		Lithostromation perdurum	
33		39		50		62		Scyphospaera spp.	
34		40		51		63		Sphenolithus abies/neoabies	
35		41		52		64		S. heteromorphus	
36		42		53		65		S. moriformis	
37		43		54		66		Sphenolithus spp.	
38		44		55		67		Reticulofenestra minuta	
39		45		56		68		R. minutula	
40		46		57		69		small R. pseudoubilicus	
41		47		58		70		R. pseudoubilicus	
42		48		59		71		"large" R. pseudoubilicus	
43		49		60		72		Rhabdosphaera spp.	
44		50		61		73		Triquetrorhabdulus milowii	
45		51		62		74		REWORKING	

Tab. 1 - Calcareous nannofossil range chart of the Tufillo section. Abundances are tabulated as percentage relative to 500 specimens. Ceratoliths and discoasters are reported as N. of specimens/mm<sup>2</sup> relative to about 10,000 coccoliths. *Discoaster* spp. are relative to 500 specimens. X = occurrence recorded out of the counting method. Letters indicate semiquantitative estimate. Total Abundance: A = 20 specimens/field of view; C = 10-20 specimens/field of view; F = 2-10 specimens/field of view; R = 0.1-1 specimen/field of view; B = barren. Relative abundance: VA = 10 specimens/field of view; A = 2-10 specimens/field of view; C = 0.5-1 specimens/field of view; F = 0.02-0.5 specimens/field of view; R = 0.005-0.02. Preservation: VP = very poor; P = poor; M = moderate.

ly on the quartzarenites of the Numidian Flysch; in the upper part it is mainly represented by marly deposits (Marne argillose del Toppo Capuana Formation).

The calcareous nannofossil assemblage allow to recognise an interval between Zone MNN3b and *A. primus* Zone (Tab. 1); a covered interval occurs between MNN7 and the *C. pelagicus* Zone. Quantitative analyses were performed from sample 7 onwards and the abundance patterns of the most significant species allow the recognition of different biozonal events (Fig. 3).

In particular the PE of small *R. pseudoumbilicus*, which represents a new event recognised in this study, is visible between samples 17-19 and it is coincident, in this section, with the FCO of *H. walbersdorfensis*. A rise in abundance of *C. leptoporus* is observed at the MNN6b/7 zonal boundary. The LO of *H. walbersdorfensis*, which indicates the top of the middle Miocene Zone MNN7, was replaced by the LCO of the species since it has been recorded up to the *Amaurolithus primus* Zone; these occurrences might be considered due to

reworking, but rare occurrences of *H. walbersdorfensis* can also be noted in the Monte dei Corvi and Contessa Entellina sections (Fornaciari et al., 1996, figs. 10-11) and in the Monte Cantigaglione section (Sprovieri et al., 1996a, fig. 6). This may suggest that the top of MNN7 Zone can be better recognised by the LCO of *H. walbersdorfensis* (Fig. 2). *Amaurolithus* spp. are extremely rare within *A. primus* Zone and were encountered in a view of 10,000 specimens. The number of reworked specimens is often moderate (<1% on 500 nannofossils) and rarely reaches more than 10% of the assemblage.

Monte Sidone section.

This composite section was sampled near Castelluccio Valmaggiore from Serra Pizzuta to Monte Cornacchia (163 III SE, topographic map of Italy); it mainly consists of alternated calcilitites, calcarenites and marls referable to the Flysch di Faeto Formation, and in the upper part of marly deposits (Marne argillose del Toppo

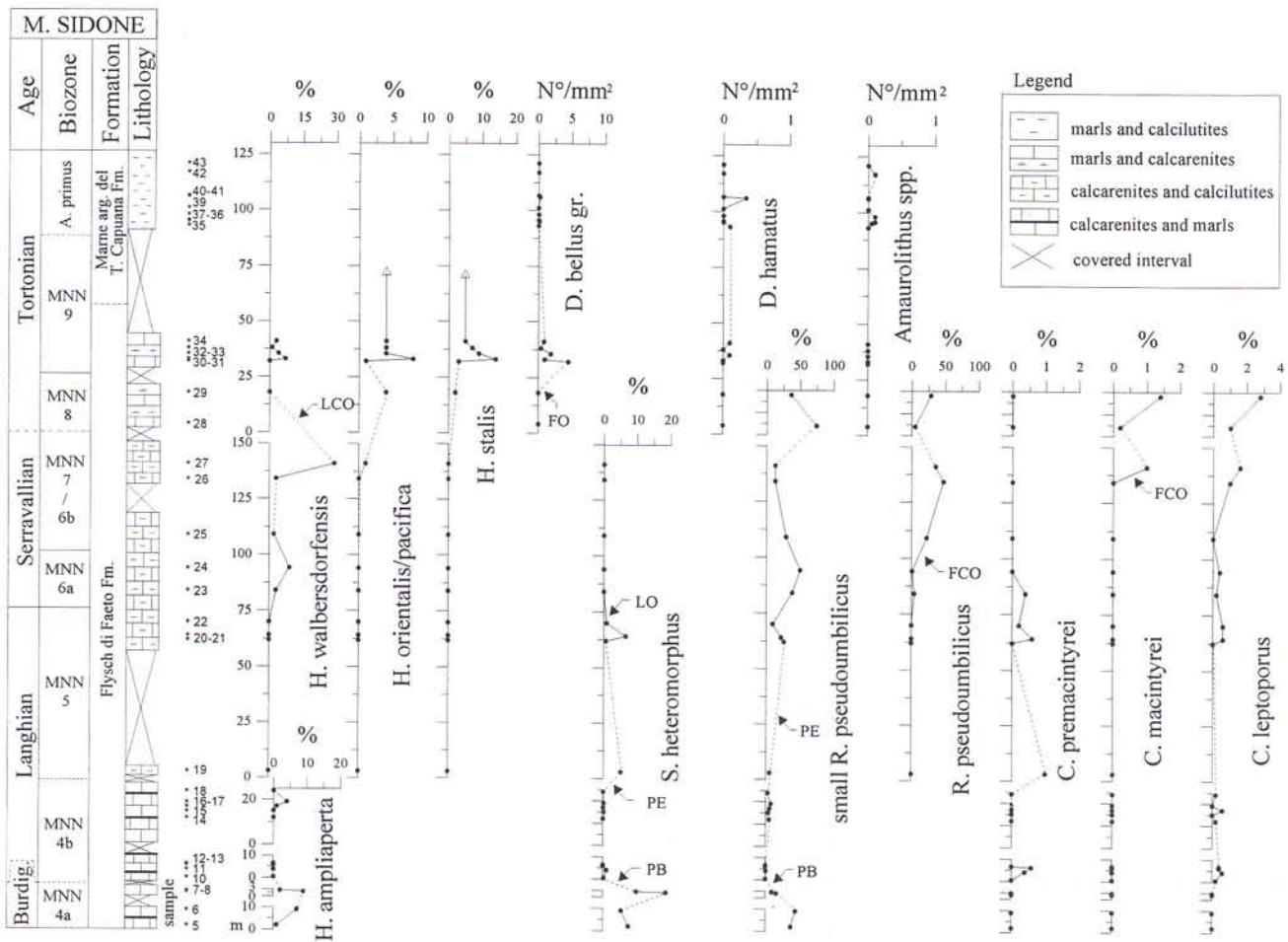


Fig. 4 - Abundance patterns of selected calcareous nannofossils at the Monte Sidone section. See legend of Fig. 3.

Tab. 2 - Calcareous nannofossil range chart of the Monte Sidone section. Abundance are reported as percentage relative to 500 specimens. From sample 30 onwards the abundance of discoasterids and ceratolithids is tabulated as N° of specimens/mm² and relative to about 10,000 coccoliths. Letters indicate semiquantitative estimate. See legend of Tab. 1.



Capuana Formation). It lies stratigraphically on the Calcarenites, marne ed argille di Monte Sidone Formation (Senatore, 1988), Aquitanian-Burdigalian in age (Russo & Senatore, 1989).

Calcareous nanofossils observed in 38 samples (Tab. 2) were referred to an interval between MNN4a and the *A. primus* Zone; an unrecovered biozonal interval occurs between MNN9 and the *A. primus* Zone. In Fig. 4 the abundance fluctuations of the most significant species can be observed. The PB and PE of small *R. pseudoumbilicus* are detectable in this section; the PB is coincident with the PB of *S. heteromorphus* and the PE is located between the PE of *S. heteromorphus* and the LO of *S. heteromorphus*. On the other hand the FCO of *H. walbersdorfensis* is not identified since the species is not recorded in the interval just below the LO of *S. heteromorphus*. The PE of small *R. pseudoumbilicus* can be useful in this case to subdivide MNN5; the LCO of *C. premacintyreii* is not a significant event: the species is rare in the section and the last occurrences are recorded below the FCO of *R. pseudoumbilicus*. The FCO of *C. macintyreii* is an alternative event to the LCO of *C. premacintyreii*. As also remarked in the Tuffillo section a rise in abundance of *C. leptoporus* occurs with the FCO of *C. macintyreii* and the LO of *H. walbersdorfensis* was replaced by the LCO of the species. Reworked specimens of *Discoaster hamatus* have been recorded within the *A. primus* Zone.

Monte Rotondo section.

The section was sampled in the neighbourhood of Bovino (174 I SE, topographic map of Italy), near Monte Rotondo and Monte Castro locality; it mainly consists of a lower and middle part of calcarenites, calcilitites and marls (Flysch di Faeto Formation) and of an upper marly interval (Marne argillose del Toppo Capuana Formation).

The total assemblage observed in the 18 collected samples is reported in Tab. 3; this allowed to refer the investigated section to an interval between MNN5 and *C. pelagicus* Zone; a stratigraphic discontinuity occurs between samples 12-13. The main quantitative results are reported in Fig. 5. The FCO of *H. walbersdorfensis* is not useful in this section to subdivide MNN5 since the species was observed from sample 12 onward only. Small specimens of *Gephyrocapsa* spp. (about 2 microns in size) have been observed within the *M. convallis* Zone (Tab. 3). It is noteworthy that specimens of small *Gephyrocapsa* have been previously recorded in the upper Miocene Mediterranean record by Bonci et al. (1991) within the Zones NN11-NN12 of Martini (1971), by Pujos (1987) in the equatorial Pacific Ocean and by Gartner (1992) in the North Atlantic within the range of *M. convallis*.

Rare and scattered reworked specimens have been recorded in this section.

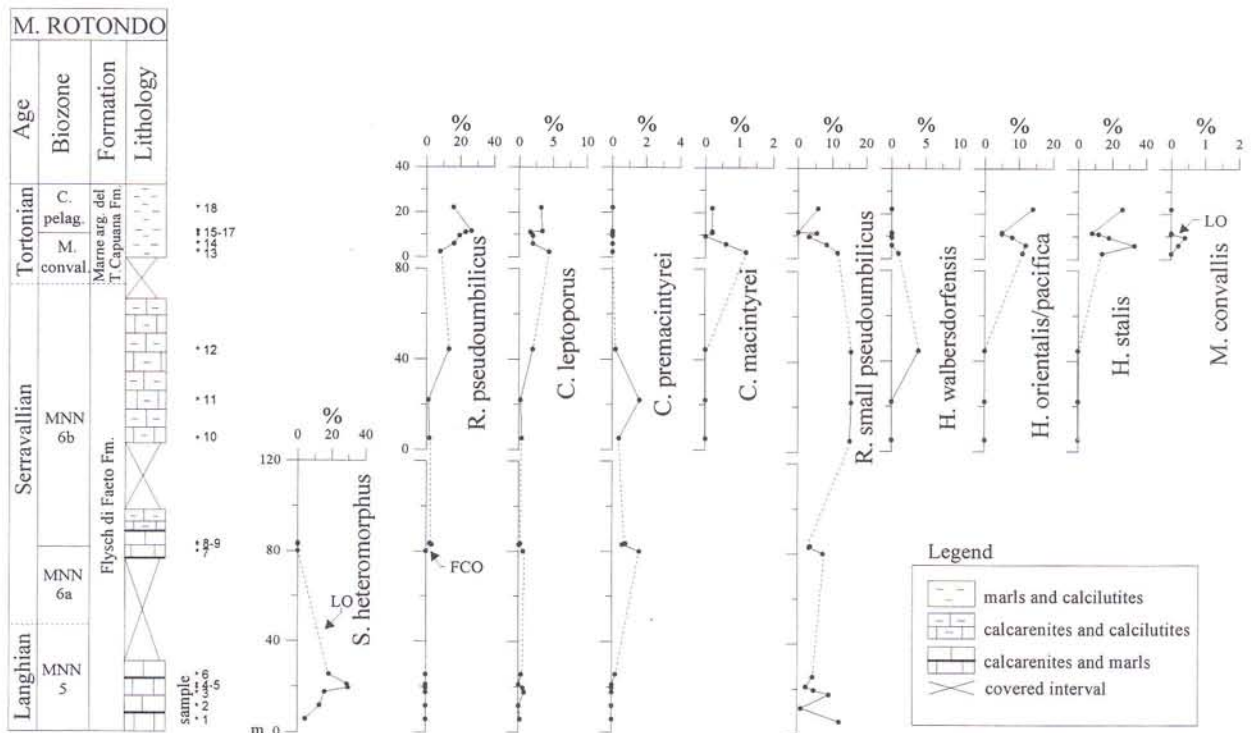


Fig. 5 - Abundance patterns of selected calcareous nanofossils at the Monte Rotondo section. See legend of Fig. 3.



Langhian		Serravallia		Tortonian		Age		
MNN5		MNN6a		M. convallis *		Zone		
Flyscht di Faeto Fm.				Marme argillose del T. Capuana Fm.		Fornaciari et al. (1996) *Theodoridis (1984)		
1	2	3	4	5	6	7	8	
C	F	C	F	F	A	C	F	
VP	VP	VP	P	P	M	M	VP	
VP	VP	VP	P	P	M	M	VP	
								Brarudosphaera bigelowii
								Calcidiscus fuscus
								C. leptoporus
								C. macintyreii
								C. premacintyreii
								Coccolithus miopelagicus
								C. pelagicus
								Cryptococcolithus mediaperforatus
								Cyclicargolithus floridanus
								Dictyococcites perplexus
								D. productus
								Discoaster adamanteus
								D. bellus
								D. bergrenii
								D. brouweri
								D. challengerii
								D. deflandrei
								Discoaster cfr D. neohamatus
								D. prepentaradiatus
								Discoaster variabilis/exilis
								Discoaster cfr D. quinquaramus
								Discoaster spp.
								Geminolithella rotula
								small Gephyrocapsa
								Helicosphaera ampliapertura
								H. carteri
								H. euphratis
								H. orientalis/pacifica
								H. stalis
								H. walbersdorfensis
								Helicosphaera spp.
								Lithostromation perdurum
								Minylitha convallis
								Pontosphaera spp.
								Pyrocyclus spp.
								Reticulofenestra minuta
								R. minutula
								small R. pseudoumbilicus
								R. pseudoumbilicus
								Rhabdosphaera spp.
								Scapholithus fossilis
								Scyphosphaera spp.
								Sphenolithus abies/neoabies
								S. moriformis
								S. heteromorphus
								Syracosphaera spp.
								Tetralithoides symeonidesii
								Triquetrorhabdulus rugosus
								REWORKING

Tab. 3 - Calcareous nannofossil range chart of the Monte Rotondo section. Abundances are reported as percentage relative to 500 specimens. From sample 13 onwards the abundance of discoasterids is tabulated as N° of specimens/mm<sup>2</sup> and relative to about 10,000 coccoliths. Letters indicate semiquantitative estimate. See legend of Tab. 1.

M. ARMENIA		Age	Formation	Sample	Total abundance	Preservation	Calciscus fuscus	C. leptopus	C. macintyreii	C. premacintyreii	Coccolithus miopelagicus	C. pelagicus	Cryptococcolithus mediaperforatus	Coronocyclus nitescens	Cyclicargolithus floridanus	Dictyoceccites perplexus	D. productus	Discoaster deflandreii	D. variabilis/exilis	Discoaster spp.	Geminitithella rotula	Helicospaera ampliaperata	H. carteri	H. euphratis	H. walbersdorfensis	Pontosphaera spp.	Rhabdosphaera spp.	Reticulofenestra minuta	R. minutula	small R. pseudoumbilicus	R. pseudoumbilicus	Sphenolithus heteromorphus	S. moriformis	Sphenolithus spp.	Tetralithoides symeonidesii	REWOKING		
Serrav.	MNN6a			12	F VP	0.2				0.4	C												0.2		0.2						F	F	0	0.2				
Langhian	MNN 5b	Flysch di Faeto Fm.	11	F P				0.2	2	67.6					0.2	0.4				0.2	0.6		20		0.6					3.2	0.2		1	0.6	2.8	0.4		
			10	F M	0.6	X		X	1.8	59.8	0.2				7.8	4						1.2	5.4		5		0.2	0.2			5.4	0.6	X	6.6	0.2	0.8	0.2	
			9	F M	1.2				0.6	3	64.6	0.2			1.8	2	C					0.8	4.6		3.8		0.2	0.8		F	0.6	0.8	0.4	12.2	0.6	1.6		0.2
			8	F M	0.6	X			0.2	3.4	65.2	X			2	1.4	C					1	3.8		7.4		X		F	1.6	1.8		9.6	0.2	1.4	0.4		
			7	F P	0.2				0.2	4.6	61	X			0.2	0.6	C	0.2				0.4	3		7.4	X	0.4		R		0.8	X	18.8		1.8		0.4	
			6	F M	0.6				0.8	6.8	57.2	0.4	X		0.2	3.4						0.2	2	1.6	X	5.8		0.2				8.8		11.6	0.2	0.2		
			5	F P	X				0.8	1.4	34.6	2					36	C				1.2	1.2		6.4			0.2	F	2.8		11.2	1.2	0.4	0.6			
	4		C M	0.6	0.6	?	0.2	4.4	47	0.2				0.8	18	C					0.4	2		6.4			0.2			5		13.4	0.4	X	0.4			
	3		F M	0.6		X	2.8	55.8	0.6		6.2	2.6	C	X							0.2	0.2		16	0.2	X	0.2	X	F			14.2		0.2	0.2			
	2		F P	0.4		X	3.2	63	0.2		8.8											0.2		19.4	0.4								4.2	X	0.2			
	1		F VP	0.6		0.4	1.2	64.8			8.2	0.4										0.2	0.2	0.4	20						0.8		0.2	2.6				

Tab. 4 - Calcareous nannofossils range chart of the Monte Armenia section. Abundance are reported as percentage relative to 500 specimens. Letters indicate semiquantitative evaluation. See legend of Tab. 1.

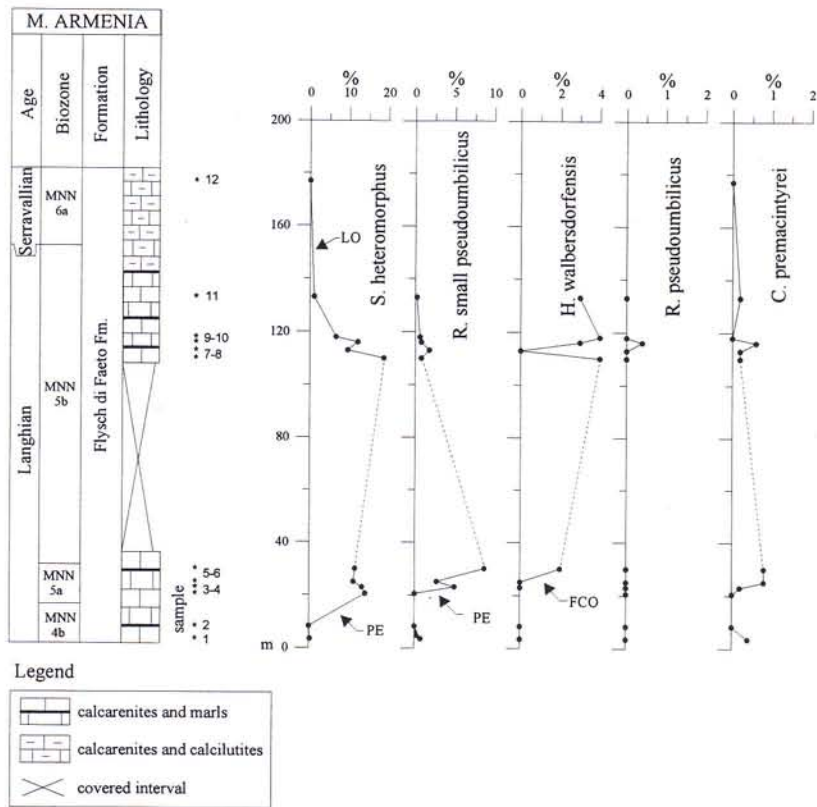


Fig. 6 - Abundance patterns of selected calcareous nannofossils at the Monte Armenia section. See legend of Fig. 3.

Monte Armenia section.

The section crops out near Forenza between Masseria Morlino and Monte Armenia (187 I SE, topographic map of Italy); it is represented by calcarenites and calcilutites with interbedded marly layers and is referable to the Flysch di Faeto Formation.

According to the total nannofossil assemblage recognised in 13 samples (Tab. 4) the section was referred to an interval between Zones MNN4b and MNN6a; the most relevant quantitative results are shown in Fig. 6.

Particularly it is visible that the PE of small *R. pseudoumbilicus* occurs between the PE of *S. heteromorphus* and the FCO of *H. walbersdorfensis*; the latter event is detectable between samples 5-6. Very rare reworked specimens occur in this section.

Fosso Scannapapere section.

The section was sampled southwest of the Forenza neighbourhood along the Fosso Scannapapere (187 I SE, topographic map of Italy).

It mainly consists of siliciclastic sandstones and marls; in the lower part of the section quartzarenite deposits, referable to the Numidian Flysch have been sampled. In the upper siliciclastic deposits (Serra Palazzo Formation) 18 samples have been considered for quantitative analyses; the complete section can be referred to an interval between Zones MNN3a and MNN5a. The total nannofossil assemblages can be found in Tab. 5 and the quantitative data of selected species are reported in Fig. 7. Common occurrences of *R. pseudoumbilicus* are observed within Zone MNN3a-3b, well below the FCO of the species which defines the top of Zone MNN6a. Moreover occurrences of *H. ampliaperta* as well as of *R. pseudoumbilicus* are recorded within the paracme interval of *S. heteromorphus* as also documented by Fornaciari et al. (1996). At the top of the section the PB of small *R. pseudoumbilicus* occurs between the PB and PE of *S. heteromorphus*. The high number of reworked specimens recorded in this section (Tab. 5) did not prevent the recognition of the biozonal events.

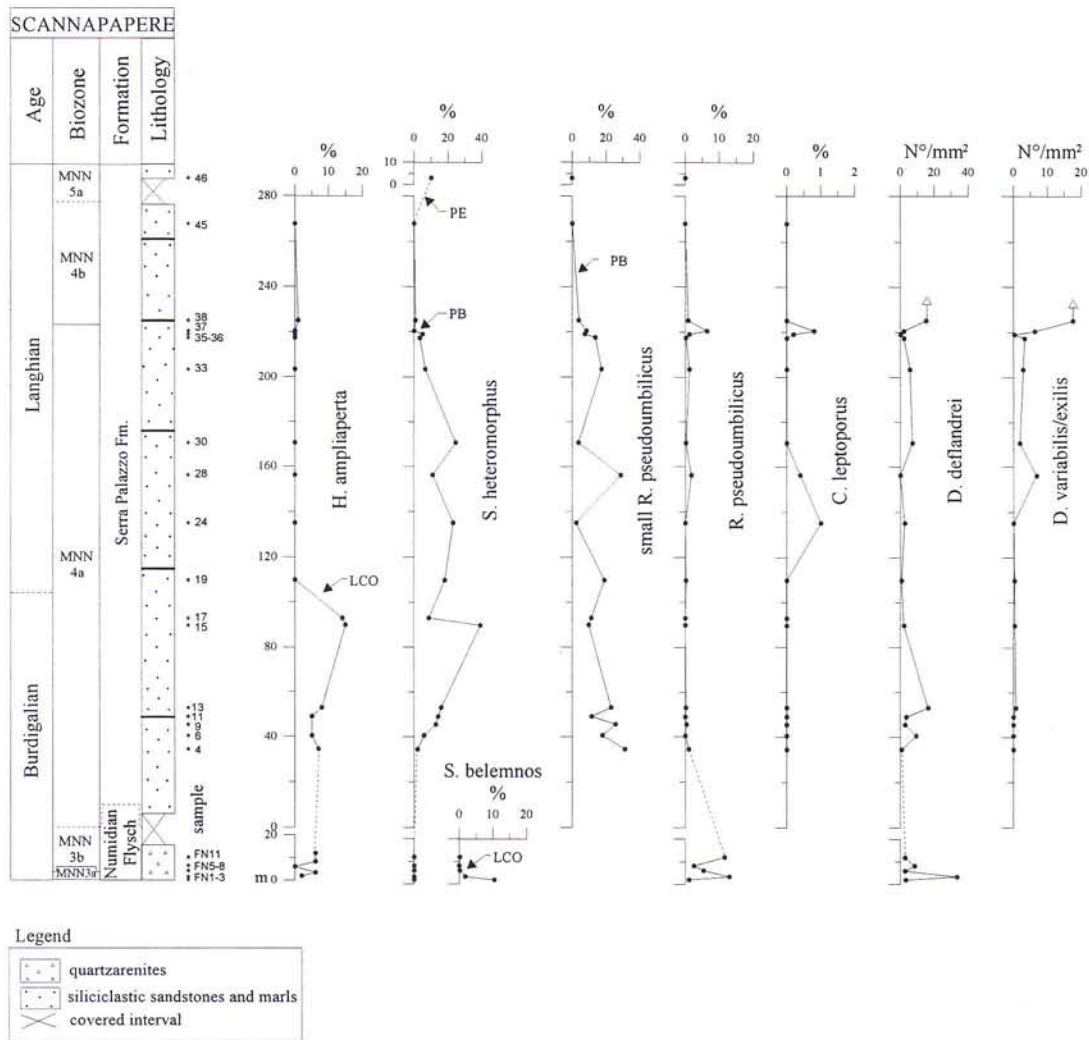


Fig. 7 - Abundance patterns of selected calcareous nannofossils at the Fosso Scannapapere section. See legend of Fig. 3.

Burdigalian		Langhian		Age		F. SCANNAPAPERRE
MNN 3b		MNN 4a		Zone		
MNN 3a				Fornaciari & Rio (1996) Fornaciari et al. (1996)		
Numidian Flysch		Serra Palazzo Fm.		Formation		
FNI		FNI		Sample		
F		F		Total Abundance		
P		P		Preservation		
						Calcidiscus fuscus
						C. leptoporus
						C. premacintyreii
						Coccolithus pelagicus
						C. miopelagicus
						Cryptococcolithus mediaperforatus
						Coronocyclus nitescens
						Cyclicargolithus floridanus
						Dictyococcites perplexus
						D. productus
						Discoaster deflandrei
						D. variabilis/exilis
						D. woodringii
						Discoaster spp.
						Geminilithella rotula
						Helicosphaera ampliaperata
						H. carteri
						H. euphratis
						H. obliqua
						H. perch-nielseniae
						H. scissura
						H. walbersdorfensis
						Helicosphaera spp.
						Pontosphaera spp.
						Pyrocyclus orangensis
						Reticulofenestra minuta
						Reticulofenestra minutula
						small R. pseudoubilicus
						R. pseudoubilicus
						Rhabdosphaera spp.
						Sphenolithus abies/neobies
						S. belemnus
						S. conicus
						S. dissimilis
						S. heteromorphus
						S. moriformis
						Sphenolithus spp.
						Tetralithoides symeonidesii
						Thoracosphaera spp.
						Triquetrorhabdulus milowii
						REWORKING

Tab. 5 - Calcareous nannofossils range chart of the Fosso Scannapapere section. Abundance are reported as percentage relative to 500 specimens. Letters indicate semiquantitative evaluation. See legend of Tab. 1.

Torrente Vaggianello section.

The section was sampled southwest of Oppido lucano, along the Vaggianello stream (188 III SW, topographic map of Italy). It mainly consists of siliciclastic sandstones in the lower part and of alternating lime-

stones, marls and silts in the upper part; it can be referred to the Serra Palazzo Formation.

The nannofossil assemblages observed in 16 samples is documented in Tab. 6; the most significant quantitative results are reported in Fig. 8. The section can be

T. VAGGIANELLO																																			
Age	Zone Fornaciari et al. (1996)	Formation	Sample	Total Abundance	Preservation	Calcidiscus leptoporus	C. macintyreii	C. premacintyreii	Coccolithus miopelagicus	C. pelagicus	Cryptococcolithus mediaperforatus	Cyclicargolithus floridanus	Dietycocites perplexus	D. productus	Discoaster deflandrei	D. variabilis/exilis	Discoaster spp.	Geminitithella rotula	Helicosphaera ampliapertura	H. carteri	H. euphratis	H. walbersdorfensis	H. waltrans	Helicosphaera spp.	Reticulofenestra minutula	small R. pseudoumbilicus	R. pseudoumbilicus	Rhabdosphaera spp.	Scapholithus fossilis	Sphenolithus abies/neoabies	S. heteromorphus	Sphenolithus moriformis	Sphenolithus spp.	REWORKING	
						Serravallian	MNN 7	Serra Palazzo Fm.	16	R	VP	R		X	F											R						R			
			15	F	VP	F	X		X	F		C								X						R	X								F
			14	F	VP	1	X	0.2	0.8	F	0.2	0.2	F				0.2		0.4						F	R	1		0.2					5.8	
			13	F	VP	3	0.2		1.4	C	X	1		A		0.4			5.8	0.8	0.4				1.8	C	C	1.6	X		0.6			21.2	
			12	F	VP		0.2	0.2	3.6	46	0.2	0.4	10.4				0.4		6.2	0.6				1	8.2	3.6	3		0.6	0.2			15.2		
			11	F	VP	2.2	0.2		0.6	53	0.2	0.2	18				0.2	0.2	11.8		0.6			1	5.4	2.2	1.4		0.4				2.6		
			10	F	P	2.6	0.2	0.4	0.2	C		0.2	C	C		0.2		0.2	13	0.2	0.2			0.2	C	F	1.2		0.2	1			20		
			9	F	P	1			0.2	49.6	0.4		11.2			0.4	0.4	10.2		0.6				3.4	4	12.8	0.2			X			5.6		
			8	F	P	0.8			0.8	40.2	0.8	0.4	3.2	C		0.6	0.4	0.4	10.2	0.4	0.2			0.2	18.8	9.2	2.2		1	0.2			10		
		MNN 6b	7	F	P	0.6		3	2	54		0.6				1		5.6							12.2	19.8							1.2		
			6	F	P			0.8	0.8	C		F	F					5.8		0.2				C	F	6		1					3.8		
			5	F	P	1		0.6	0.6	C	0.2	F	R	1.6		1	0.2	5.6	0.2	X	0.2	C	F	5.8		C	F	5.8		0.2	0.2	0.2	6.2		
		MNN 4a	4	F	M			0.2	0.2	37		1.2						0.2	11.2					3.8	5.2					40.8			0.2		
			3	F	P					52.4		1.6						0.4	14					4.6	15.2				11.2	0.2			0.4		
			2	F	VP					16.4		2.4		0.4	5.2			1.2	3.2					24.4	10.2	1.6		0.6	11.6	0.4		22.4			
			1	F	VP				0.2	11.4		2.8		1.2	3			1.6	1.6					26.2	19.6	3.6		1.4	11.6			15.8			

Tab. 6 - Calcareous nannofossil range chart of the Torrente Vaggianello section. Abundance are reported as percentage relative to 500 specimens. Letters indicate semiquantitative estimate. See legend of Tab. 1.

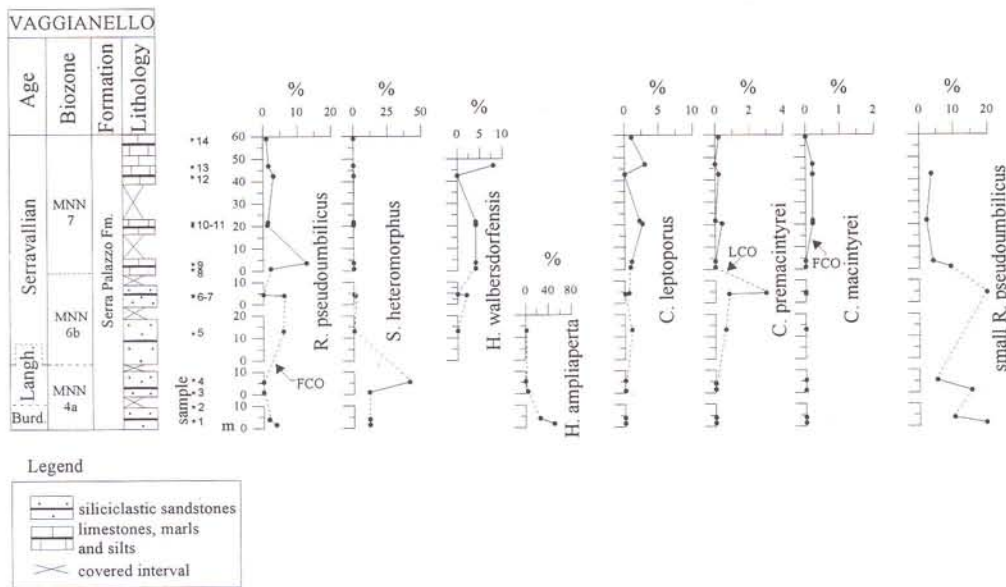


Fig. 8 - Abundance patterns of selected calcareous nannofossils at the Torrente Vaggianello section. See legend of Fig. 3.

referred to an interval between MNN4a and MNN7 with a major stratigraphic discontinuity between samples 4-5. In the lower part of the section the presence of *S. heteromorphus*, *H. ampliaperta* and small *R. pseudoumbilicus* represent MNN4a; *R. pseudoumbilicus* is recorded from MNN4a upward. The LCO of *C. premacintyreii* as well as the FCO of *C. macintyreii* can be observed in this section. Reworked specimens were found with percentages that varie from 0.2 up to 20% of the assemblage.

Jazzo Porcellini section.

The section was sampled near Stigliano along the Vallone della Difesa, at the Jazzo Porcellini locality (200 II NW, topographic map of Italy). In the lower and middle part it is represented by silicilastic sandstones overlying the Numidian Flysch (Gallicchio, 1996) and by limestones, marls and silts referable to the Serra Palazzo Formation. In the upper part it mainly consists of marly deposits referable to the Marne argillose del Toppo Capuana Formation.

The nannofossil assemblages observed in 30 samples are reported in Tab. 7; the quantitative results of the significant species are shown in Fig. 9. The Jazzo Porcellini section has been referred to an interval between Zone MNN4a and Zone MNN8.

In sample 9, within the paracme interval, an abundance peak of *S. heteromorphus*, probably due to reworking, is recorded. Several Cretaceous and Paleogene specimens occur in the same sample (Tab. 7). However, it is noteworthy that also Fornaciari et al. (1996) recorded a characteristic abundance peak of *S. heteromorphus* within the paracme interval of the species and correlatable among a few sections.

The FCO of *H. walbersdorfensis* as well as the LCO of *C. premacintyreii* were not detected in this section, probably due to stratigraphic discontinuities. The LCO of *H. walbersdorfensis* may be doubtfully placed at the top of the section, between samples 29-30. The number of reworked specimens is considerable in this section and reaches values up to 40% of the assemblage.

Valle Difesa section.

The section was sampled near Stigliano at the Piceca locality (200 II NW, topographic map of Italy); it mainly consists of siliciclastic sandstones and of alternating limestones, marls and silts in the uppermost part and is referable to the Serra Palazzo Formation. The biostratigraphic study performed on 21 samples allows to recognise an interval between Zones MNN4a and

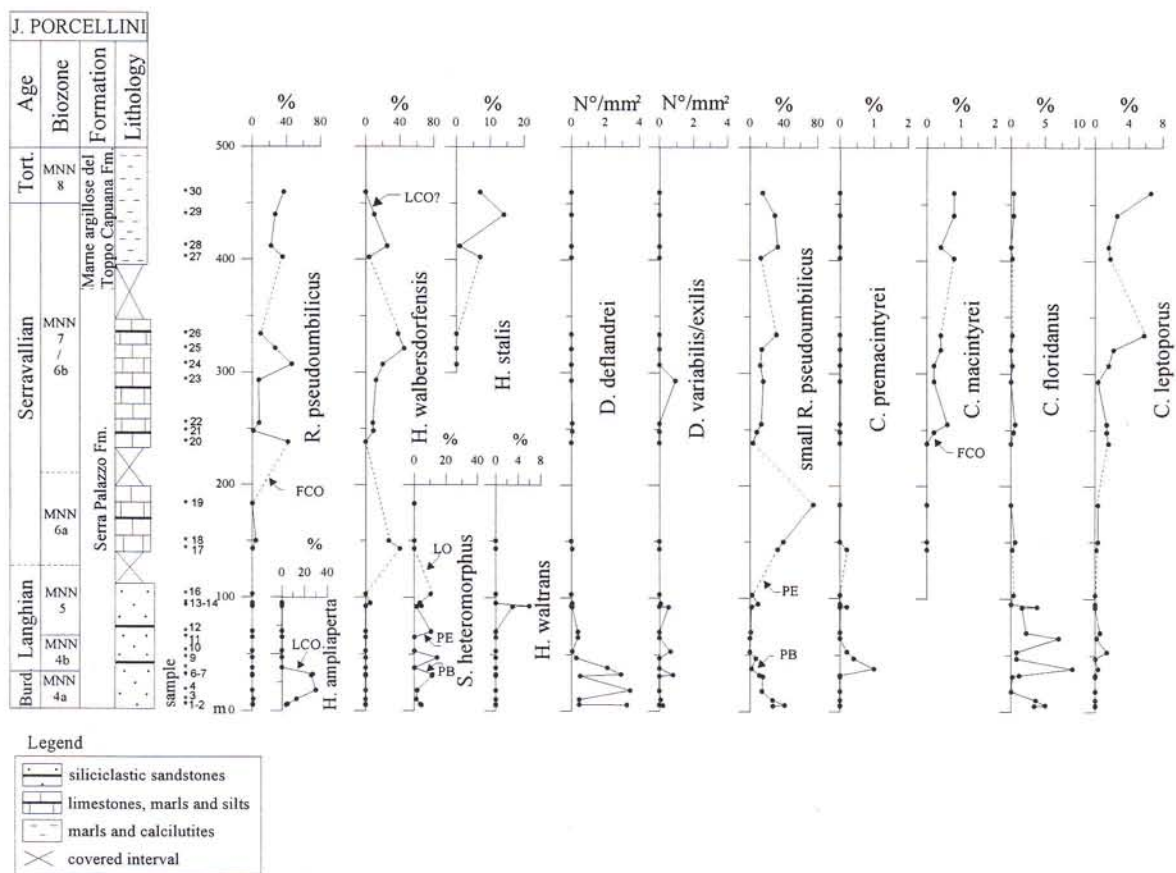


Fig. 9 - Abundance patterns of selected calcareous nannofossils at the Jazzo Porcellini section. See legend of Fig. 3.

Burdigalian		Langhian		Serravallian				Torton.	Age		Zone	JAZZO PORCELLINI
MNN 4a		MNN 4b		MNN 5		MNN 6a		MNN 6b / 7		MNN 8		
Serra Palazzo Fm.								M. arg. del T. Cap. Fm.	Formation			
1	C	2	C	3	F	4	F	5	M	6	C	Sample
	M		M		P		P		M		M	Total Abundance
0.2												Preservation
												Calcidiscus fuscus
												C. leptoporus
												C. macintyreii
												C. premacintyreii
												Cryptococcolithus mediaperforatus
												Coccolithus miopelagicus
												C. pelagicus
												Cyclcargolithus floridanus
												Geminilithella rotula
												Dictyococcites perplexus
												D. productus
												Discoaster deflandrei
												D. exilis
												D. moorei
												D. musicus
												D. variabilis/exilis
												D. variabilis
												Discoaster spp.
												Helicosphaera ampliaperpta
												H. carteri
												H. euphratis
												H. orientalis
												H. perch-nielseniae
												H. scissura
												H. stalis
												H. walbersdorfensis
												H. waltrans
												Helicosphaera spp.
												Lithostromation perdurum
												Pyrocyclus inversus
												Reticulofenestra minuta
												R. minutula
												small R. pseudoubilicus
												R. pseudoubilicus
												Rhabdosphaera spp.
												Sphenolithus abies/neoabies
												S. heteromorphus
												S. moriformis
												Sphenolithus spp.
												Syracosphaera spp.
												Tetralithoides symeonidesii
												Triquetrorhabdulus rugosus
												REWORKING

Tab. 7 - Calcareous nannofossil range chart of the Jazzo Porcellini section. Abundance are reported as percentage relative to 500 specimens. Letters indicate semiquantitative estimate. See legend of Tab. 1.

MNN6b/7 (Tab. 8). The quantitative patterns of selected species are shown in Fig. 10.

The FCO of *H. walbersdorfensis* and the LCO of *C. premacintyreii* are not usable in this section to subdivide Zone MNN5 and MNN6b/7 respectively; on

the other hand the FCO of *C. macintyreii* provides a useful event within the MNN6b/7 interval and is associated with an increase in abundance of *C. leptoporus*. Numerous reworked specimens were found in this section with percentages that may reach values up to 30%.

Langhian		Serravallian										Age											
MNN 4a	MNN 4b	MNN 5					MNN 6a					MNN 7 / 6b	Zone	Fornaciari et al. (1996)									
Serra Palazzo Fm.												Formation											
1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18	19	20	21	Sample	Total Abundance	Preservation
C	F	F	F	C	F	F	F	F	F	F	F	F	F	C	C	F	F	F	F	F			
X																							Calcidiscus fuscus
																							<i>C. leptoporus</i>
																							<i>C. macintyreii</i>
																							<i>C. premacintyreii</i>
																							<i>Coccolithus miopelagicus</i>
																							<i>C. pelagicus</i>
																							<i>Cryptococcolithus mediaperforatus</i>
																							<i>Cyclicargolithus floridanus</i>
																							<i>Geminitrileta rotula</i>
																							<i>Dictyococcites perplexus</i>
																							<i>D. productus</i>
																							<i>Discoaster deflandrei</i>
																							<i>D. exilis</i>
																							<i>D. variabilis/exilis</i>
																							<i>D. variabilis</i>
																							<i>Discoaster</i> spp.
																							<i>D. moorei</i>
																							<i>D. signus</i>
																							<i>D. musicus</i>
																							<i>Helicosphaera ampliaptera</i>
																							<i>H. carteri</i>
																							<i>H. euphratis</i>
																							<i>H. obliqua</i>
																							<i>H. orientalis</i>
																							<i>Helicosphaera stalis</i>
																							<i>H. walbersdorfensis</i>
																							<i>H. waltrans</i>
																							<i>Helicosphaera</i> spp.
																							<i>Rhabdosphaera</i> spp.
																							<i>Reticulofenestra minuta</i>
																							<i>R. minutula</i>
																							<i>R. pseudoubilicus</i>
																							small <i>R. pseudoubilicus</i>
																							<i>Sphenolithus abies/neoabies</i>
																							<i>S. heteromorphus</i>
																							<i>S. moriformis</i>
																							<i>Sphenolithus</i> spp.
																							<i>Syracosphaera</i> spp.
																							<i>Tetralithoides symeonidesii</i>
																							<i>Triquetrorhabdulus rugosus</i>
																							REWORKING

Tab. 8 - Calcareous nannofossil range chart of the Valle Difesa section. Abundance are reported as percentage relative to 500 specimens. Letters indicate semiquantitative estimate. See legend of Tab. 1.



Rotondella section.

The section was sampled in the neighbourhood of Rotondella and along the Canale Ruggero (212 III NW, topographic map of Italy); the section, referable to the Flysch di Faeto Formation, mainly consists of calcarenites, calcilutites and marls in the lower part and of limestones, marls and silts upwards.

Nannofossils are rare and poorly preserved in this section; quantitative distribution of 21 samples was collected by counting the most significant species in 300 fields of view. Results are reported in Tab. 9 and Fig. 11.

The LO of *S. heteromorphus* between samples 3-4 was recognised; the common occurrences of *R. pseudoubilicus* and of *C. macintyreii* allow to recognise the MNN6b/7 interval. The abundance pattern of *C. premacintyreii* is not useful in this section.

DSDP Leg 42, Site 372 - western Mediterranean Sea.

Site 372, which represents a reference section for the lower and middle Miocene biostratigraphy in the Mediterranean region, is located on the Menorca Rise of the Balearic Basin (40°01.86' N, 04°47.79' E). The main quantitative results are shown in Fig. 12 and may be summarized as follows:

a) The abundance pattern of small *R. pseudoubilicus* shows a distinct paracme interval within Zones MNN4b and 5a; the PB occurs shortly above the PB of *S. heteromorphus* and the PE slightly below the FCO of *H. walbersdorfensis*. This is well comparable with results obtained in the on-land sections confirming the biostratigraphic value of the newly recognised events;

b) The distribution of *R. pseudoubilicus* is similar to the one observed in the Fosso Scannapapere section: the species is recorded well below its FCO and was noted within Zones MNN 2b-3a-3b-4b.

Moreover, in the lower part of the section the FO of *H. mediterranea* was recorded below the FO of *H. ampliapertura*, within Zone MNN2a (core 39). On the contrary Fornaciari et al. (1996) documented the FO of *H. mediterranea* above the FO of *H. ampliapertura* in the Mediterranean region. It is possible that the presence of several unrecovered intervals at Site 372 prevented the recognition of *H. ampliapertura* below core 39. However, according to Theodoridis (1984) and Perch-Nielsen (1985) the occurrences of *H. mediterranea* precedes that of *H. ampliapertura*; moreover recently de Kaenel & Villa (1996) recorded at the Iberia Abissal Plain (ODP Leg 149) the FO of *H. mediterranea* below the FO of *H. ampliapertura*.

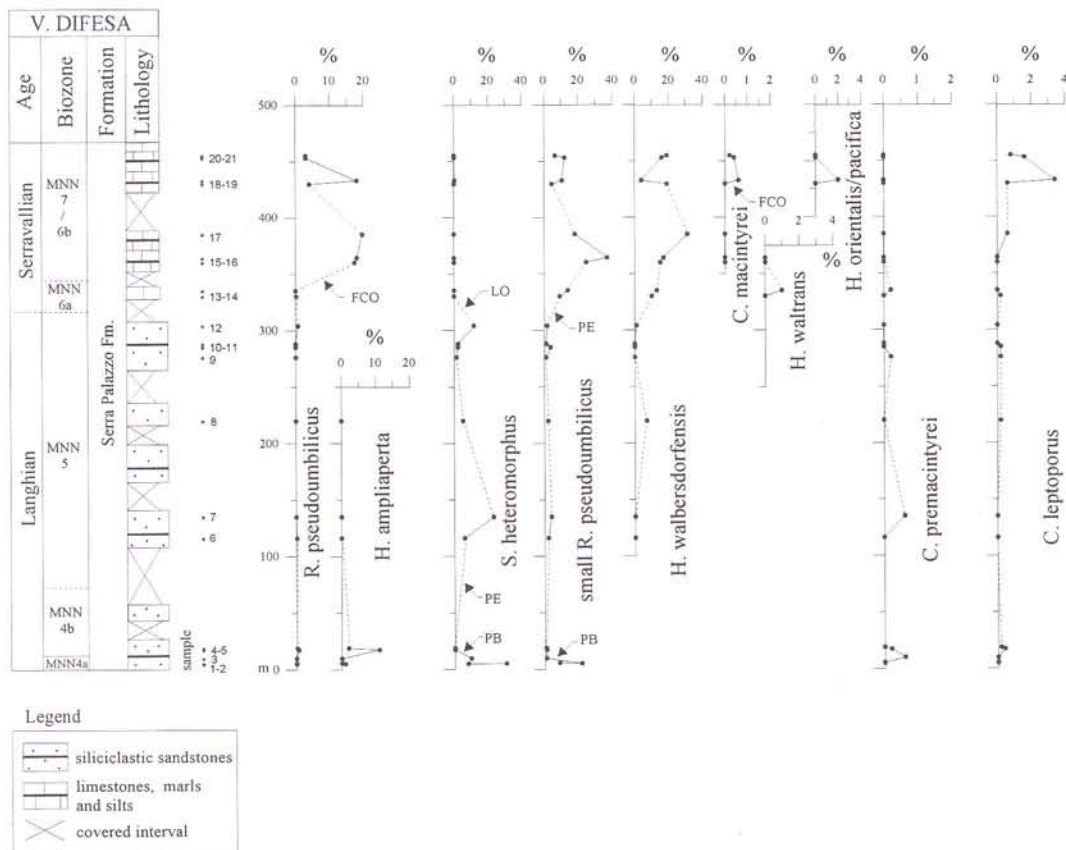


Fig. 10 - Abundance patterns of selected calcareous nannofossils at the Valle Difesa section. See legend of Fig. 3.

ROTONDELLA		Age		Zone		Formation		Sample		Total abundance		Preservation		Fossil List			
		Formiaci et al. (1996)															
Serravallian	MNN 7 / 6b	Flysch di Faeto Fm.	19	F	VP	15	3.8	0.6	C								
			18	F	VP	2.4	0.2		23.4	0.6	11.5	F		0.2	1.7		
			17	F	P	1.7	0.4				0.6	R	F		0.2		
			16	F	VP	3.6	0.4	0.2	F			R	R		0.6		
			15	F	P	2.1			27.2	0.4	9.2	F			3.4		
			14	F	VP	1.1	0.4	0.2	22.5	0.2	1.6				1.7		
			13	F	P	5	0.2	0.2	A	0.4	C	C	0.2		3.4	0.2	0.2
			12	F	P	5.8	1.5	0.2	A	0.2	0.4	F	F		3	0.2	0.2
			11	F	R	8	0.8		41.2			A			1.6		
			10	F	VP	15.5	0.7	0.9	C			F	R		0.4		
			9	F	P	8.1	2.3	0.7	A			A			1.5		
			8	F	P	7.1	0.2	0.7	C	0.2		A			0.4		
			7	F	VP	5	0.9	0.9	C		0.2	R			0.7		
			6	F	P	20.2	0.2	0.6	3	A		C	F	1.1	14.3	0.4	
			5	F	VP	11.4	1.3	2.3	C		0.6	F	F		2.3		
			4	F	VP	2.6	0.2	2.8	C		0.6	C	C		3.6		
			3	F	P	0.4	0.4		C		17.4	R	1.2	0.2	14.6		
			2	F	VP	1.9			C		3.6				6.8		
			1	F	VP	1.1	2.9	14.2	A		R		1.1		F		

Tab. 9 - Calcareous nannofossil range chart of the Rotondella section. Abundance are reported as N° of specimens/mm<sup>2</sup> and relative to 300 fields of view. Letters indicate semiquantitative estimate. See legend of Tab. 1.

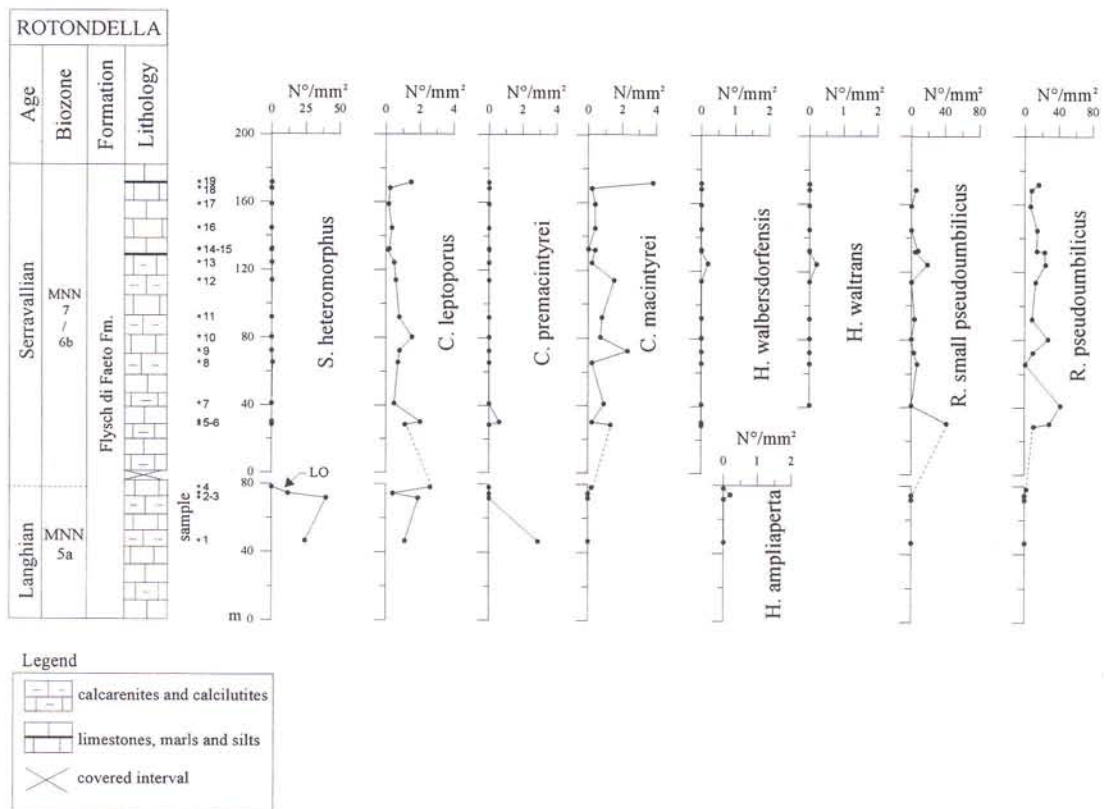


Fig. 11 - Abundance patterns of selected calcareous nannofossils at the Rotondella section. See legend of Fig. 3.

**Comparison with previous age assignments in the Southern Apennines sections.**

The previous biostratigraphic data on the studied deposits (Serra Palazzo and Flysch di Faeto formations) are mostly based on planktonic foraminiferal assemblages; the comparison between the results obtained in this study and the previous ones thus is not simple, mainly because these latter refer to biostratigraphic and chronostratigraphic terminologies that can be considered out of use at present. Moreover, the correlation between the calcareous nannofossil scheme adopted in this study and the foraminiferal biostratigraphy is not completely established.

For these reasons a synthesis of the main foraminiferal biostratigraphic results obtained in the foredeep sections by previous authors is shown in Tab. 10. In Tab. 11 a tentative correlation between the past chronostratigraphic units adopted in the studied turbidite deposits and the chronostratigraphy used in this study is presented.

The obtained biostratigraphic results provide, for the Serra Palazzo and Flysch di Faeto formations, an age close to the Burdigalian/Langhian boundary (Zones MNN4a-4b) and the Serravallian-Tortonian (Zones MNN7-9) following the chronostratigraphy proposed by

Fornaciari et al. (1996). In the upper marly portion of the studied sections (Marne argillose del Toppo Capuana Formation) the interval between MNN7 and *A. primus* Zone has been found.

Previous studies often referred the studied deposits to the upper Langhian - Tortonian (Tab. 10); the discrepancies with the present results mainly concern the age assigned to the lower part of the investigated deposits and this is mainly due to differences in the chronostratigraphic terminology (Tab. 11).

On the other hand, a slightly younger biozonal assignment was inferred by Palmentola (1970) that recognised the *O. suturalis* subzone since the lower part of Serra Palazzo Formation, as well as by Patacca et al. (1991) which recorded *O. suturalis* and *O. bilobata* from the lower part of the Tuffillo and Flysch di Faeto formations. Patacca et al. (1992) also suggested a Langhian age for the "Apennine Numidian quartzarenites" (stratigraphically underlying the studied deposits); however they did not record the presence of *Praeorbulina glomerosa sicana* within the "Numidian interval" (tab. 3, p. 323) but referred to data from Carbone et al. (1987).

For the moment the conflicting data in the lower part of the studied sections cannot be interpreted sati-

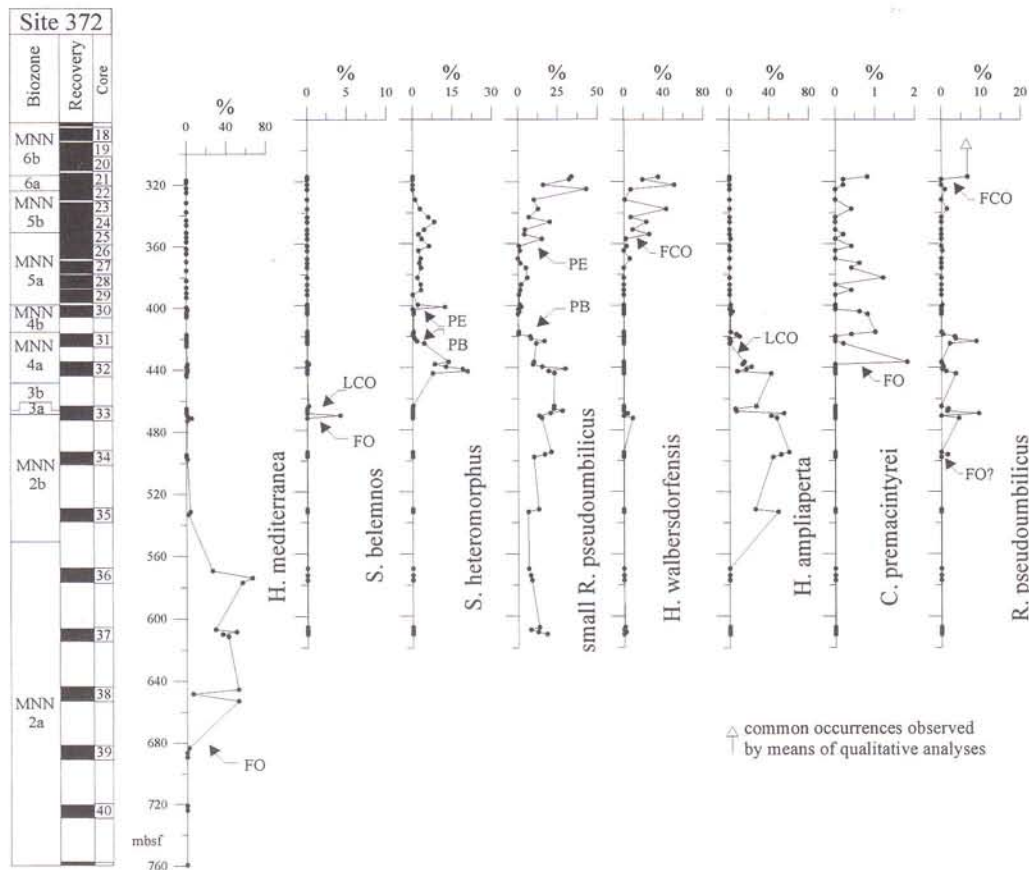


Fig. 12 - Abundance patterns of selected calcareous nannofossils at DSDP Site 372. See legend of Fig. 3.

Author	Formation	Biostratigraphic reference	Biozone	Age
Selli (1962)	Serra Palazzo F.	Selli (1957)		late Langhian-early Helvetian
Selli (1962)	Tufillo F.	Selli (1957)		late Langhian - early Helvetian
Crostella & Vezzani (1964)	Flysch di Faeto F.			Helvetian-early Tortonian
Casnedi (1964)	Serra Palazzo F.	Selli (1957)	lower part of <i>Orbulina universa</i> Zone	early Helvetian
Wezel (1966)	Serra Palazzo F.	Wezel (1966)	<i>Globorotalia fohsi</i> Zone	early-middle Helvetian
Palmentola et al. (1967)	Stigliano F./Serra Palazzo F. boundary	Crescenti (1966)	<i>Globigerinoides trilobus</i> Zone	Langhian
Pieri & Radina (1967)	Stigliano F./Serra Palazzo F. boundary	Crescenti (1966)	<i>Globigerinoides trilobus</i> Zone	Langhian
Boenzi et al. (1968)	lower part of Serra Palazzo F.	Crescenti (1966)	<i>Globigerinoides trilobus</i> Zone	Langhian
Palmentola (1970)	Serra Palazzo F.	Cati et al. (1968)	<i>Orbulina</i> s.l. Zone, subzone <i>O. suturalis</i> - <i>Globoquadrina altispira</i> / <i>Globorotalia miozea</i>	late Langhian-Serravallian
Palmentola (1970)	Flysch di Faeto F.	Cati et al. (1968)	<i>Orbulina</i> s.l. - <i>Globorotalia menardii</i> Zones, subzone <i>Globorotalia ventriosa</i> and <i>Globorotalia nepenthes</i>	middle-late Serravallian /early-middle(?)Tortonian
Ciaranfi et al. (1973)	Serra Palazzo F.	Cati et al. (1968)	<i>Orbulina</i> s.l. Zone, subzone (?) <i>Globoquadrina altispira</i>	Serravallian
Pieri & Walsh (1973)	Serra Palazzo F.	Cati et al. (1968)	<i>Globigerinoides trilobus</i> Zone, subzone <i>Globigerinoides bisphericus</i> and in the upper part <i>Praeorbulina glomerosa</i> s.l. Zone - <i>Orbulina</i> s.l. Zone.	middle Langhian - early Serravallian
Rapisardi & Walsh (1978)	Flysch di Faeto F.		significant specimens: <i>G. trilobus</i> , <i>G. altispira</i> , <i>G. dehiscens</i> , <i>O. suturalis</i> , <i>O. universa</i>	Langhian - Serravallian
Di Nocera & Torre (1987)	Flysch di Faeto F.			late Langhian - Serravallian
Russo & Senatore (1989)	Flysch di Faeto F.	Iaccarino (1985)	<i>Praeorbulina glomerosa</i> s.l., - <i>Globorotalia acostaensis</i> Zones	early Langhian - early Tortonian
Russo (1988) Russo & Senatore (1989)	Marne argillose del Toppo Capuana F.	Iaccarino (1985)	<i>Globorotalia acostaensis</i> - <i>Globigerinoides extremus</i> Zones	early-late Tortonian
Patacca et al. (1991)	Tufillo F. and Flysch di Faeto F.		<i>Orbulina</i> is present in the lower part	Serravallian-Tortonian

Tab. 10 - Main previous biostratigraphic results in the investigated deposits (Flysch di Faeto Formation, Serra Palazzo Formation, Marne argillose del Toppo Capuana Formation).

syingly. Different intervals of sampling or slightly different ages in various sections can be invoked. It is unlikely that reworking phenomena prevented the recognition of "younger" assemblages in the studied sections and at the same time generated a sequence of biostratigraphic events comparable to those observed in pelagic and hemipelagic sediments by other authors, as well as in the DSDP Site 372 in this work.

**Conclusions.**

The quantitative calcareous nannofossil biostratigraphic study performed in nine Miocene Southern Apennines foredeep sections provided new information on the applicability of the recently proposed Mediterranean biozonal schemes. Several biozonal events proposed by Fornaciari & Rio (1996) and Fornaciari et al. (1996) in

the lower and middle Miocene sections and by Theodoridis (1984) in the upper Miocene interval have been found. A Burdigalian-upper Tortonian interval between MNN3a and the *Amaurolithus primus* Zone was identified and new biostratigraphic events were proposed.

Within the lower and middle Miocene record the LCO of *H. ampliapertura*, the PB of *S. heteromorphus*, the PE of *S. heteromorphus*, the LO of *S. heteromorphus* and the FCO of *R. pseudoumbilicus*, provide a good biostratigraphic resolution of the studied sections and all furnished useful events for stratigraphic correlation.

On the other hand, some of the biohorizons of the above mentioned zonal schemes, such as the FCO of *H. walbersdorfensis* and the LCO of *C. premacintyreii*, were not always identified. *H. walbersdorfensis* is often not recorded in the interval just below the LO of *S. heteromorphus* and the rare occurrences of *C. premacintyreii*.

Foraminifera										Nannofossil	
Iaccarino (1985)		Cati et al. (1968)	Wezel (1966)	Crescenti (1966)	Selli (1957)	1	2	Biozone *			
Messin.	No distinctive Zone		Not defined								
	Gl. conomiozea		Gl. miocenica								
Tortonian	Gl.oides obliquus extremus	Gl. suturae	Gl. menardii	Gl. ventriosa G. nepenthes	Gl. menardii	Gl. menardii	not studied	C. leptoporus	R. rotaria	A. primus	
		Gl.oides obliquus extremus/ Gl.oides bulloideus									
	Gl. acostaensis										
Serravallian	Gl. menardii s.l.		Orbulina s.l.	Gl. foehsi	O. universa	O. universa	MNN9 MNN8b MNN8a	M. convallis	M. convallis	M. convallis	
	Gl. siakensis	Gl. siakensis-Gl.oides obliquus obliquus									Gl.oides obliquus
		Gl.oides subquadratus									Gl. linguaensis
	O. suturalis-Gl. peripheroronda	G. altispira altispira									G. altispira
		Gl. praemenardii - Gl. peripheroronda									Gl. miozea
		O. universa									
Langh.	O. suturalis		O. suturalis								
	P. glomerosa s.l.		P. glomerosa s.l.								
Burdig.	Gl.oides trilobus		Gl.oides trilobus	Gl.oides trilobus	Gl.oides trilobus	Globoquadrina	MNN4a	MNN3b	MNN3a	MNN2b	
	G. dehiscens	Gl.oides altiapertura	Gl.oides bisphericus								
Aquit.	C. dissimilis	G. dehiscens	G. dehiscens	G. dissimilis	G. dissimilis	G. dissimilis	MNN2a	MNN2a	MNN2a	MNN2a	
		G. dehiscens dehiscens	Gl.oides altiapertura - C. dissimilis								Gl.oides altiapertura - Gl.oides trilobus
							Langhian	MNN5a			
							Langhian	MNN4b			
							Burdig.	MNN4a			
							Burdig.	MNN3b			
							Burdig.	MNN3a			
							Burdig.	MNN2b			
							Burdig.	MNN2a			
							Aquit.	MNN1d			
							Aquit.	MNN1c			

Tab. 11 - Tentative correlations between the foraminifera biostratigraphic schemes and chronostratigraphy adopted in previous age assignments in the foredeep deposits and the biostratigraphic and chronostratigraphic scheme adopted in this work. In column 1 is reported the chronostratigraphy adopted in Selli (1957) and Crescenti (1966); in column 2 the one proposed in Fornaciari & Rio (1996) and Fornaciari et al. (1996) and adopted in this work; the Tortonian-Messinian boundary has been traced according to Langereis et al. (1984). \*: Biozone of Fornaciari & Rio (1996) and Fornaciari et al. (1996) in the lower-middle Miocene and of Theodoridis (1984) in the upper Miocene.

*rei* prevented to obtain significant abundance patterns for the recognition of the LCO of this species.

This quantitative study provides new stratigraphic ranges and biostratigraphic events, which improve the lower and middle Miocene Mediterranean biostratigraphy:

- *R. pseudoumbilicus* has been recorded well below the FCO of the species, in the turbidite deposits as well as in DSDP Site 372. Discontinuous occurrences were noted within Zones MNN2b-3a-3b-4b. It is noteworthy that this is quite useful in the biostratigraphic analyses of the on-land deposits, since the presence of *R. pseudoumbilicus* may wrongly suggest a younger age of the section;

- a distinctive paracme interval of small *R. pseudoumbilicus* was identified within Zones MNN4b and MNN5a: the PB occurs slightly above the PB of *S. heteromorphus* and represents an additional event for the recognition of MNN4b. The PE occurs shortly below the FCO of *H. walbesdorfensis* and can be considered an alternative event to subdivide MNN5. The occurrence of the paracme interval in the same stratigraphic position at Site 372 confirms its potential value for stratigraphic correlation within the Mediterranean region;

- the FCO of *C. macintyreii* is well detectable in the studied sections and represents a more reliable event than the LCO of *C. premacintyreii* within the MNN6b/7 interval. An increase in abundance of *C. leptoporus* was also noted with the FCO of *C. macintyreii*;

- the LCO of *H. walbesdorfensis* appears to be a better biohorizon than the LO of the species for the recognition of the top of Zone MNN7.

At Site 372 the FO of *H. mediterranea* precedes the FO of *H. ampliapertura*.

In the upper Miocene record the *M. convallis*, *C. pelagicus* and *A. primus* Zones of Theodoridis (1984) were recognised. In the upper part of the *M. convallis* Zone abundant specimens referable to *Gephyrocapsa* spp. (about 2 microns in size) were recorded.

The quantitative biostratigraphy was helpful in the stratigraphic reconstruction of the studied deposits, although these are often affected by stratigraphic discontinuities. The recognition of several new biozonal horizons improves the stratigraphic framework of the Miocene foredeep deposits in the Southern Apennines.

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## Appendix

List of calcareous nannofossils considered in this work.

*Amaurolithus primus* (Bukry & Percival, 1971) Gartner & Bukry, 1975  
*Braarudosphaera bigelowii* (Gran & Braarud, 1935) Deflandre, 1947  
*Calcidiscus leptoporus* (Murray & Blackman, 1898) Loeblich & Tappan, 1978  
*Calcidiscus fuscus* (Backman, 1980) Janin, 1987  
*Calcidiscus macintyreii* (Bukry & Bramlette, 1969) Loeblich & Tappan, 1978  
*Calcidiscus premacintyreii* Theodoridis, 1984  
*Catinaster coalitus* Martini & Bramlette, 1963  
*Coccolithus miopelagicus* Bukry, 1971  
*Coccolithus pelagicus* (Wallich, 1877) Schiller, 1930  
*Coronocylus nitescens* (Kamptner, 1963) Bramlette & Wilcoxon, 1967  
*Cryptococcolithus mediaeporatus* (Gartner, 1992) de Kaenel & Villa, 1996  
*Cyclicargolithus floridanus* (Roth & Hay in Hay et al., 1967) Bukry, 1971  
*Dictyococcites productus* (Kamptner, 1963) Backman, 1980  
*Dictyococcites perplexus* Burns, 1975  
*Discoaster adamanteus* Bramlette & Wilcoxon, 1967  
*Discoaster asymmetricus* Gartner, 1967  
*Discoaster bellus* Bukry & Percival, 1971  
*Discoaster berggrenii* Bukry, 1971  
*Discoaster brouweri* (Tan, 1927) Bramlette & Riedel, 1954  
*Discoaster challengerii* Bramlette & Riedel, 1954

*Discoaster deflandrei* Bramlette & Riedel, 1954  
*Discoaster exilis* Martini & Bramlette, 1963  
*Discoaster hamatus* Martini & Bramlette, 1963  
*Discoaster intercalaris* Bukry, 1971  
*Discoaster moorei* Bukry, 1971  
*Discoaster musicus* Stradner, 1959  
*Discoaster neohamatus* Bukry & Bramlette, 1969  
*Discoaster neoerectus* Bukry, 1971  
*Discoaster pentaradiatus* (Tan, 1927) Bramlette & Riedel, 1954  
*Discoaster prepentaradiatus* Bukry & Percival, 1971  
*Discoaster quinqueringii* Bramlette & Riedel, 1954  
*Discoaster signus* Bukry, 1971  
*Discoaster surculus* Gartner, 1967  
*Discoaster tamalis* Kamptner, 1967  
*Discoaster triradiatus* Tan, 1927  
*Discoaster variabilis* Martini & Bramlette, 1963  
*Discoaster variabilis/exilis* (sensu Rio et al. 1990a)  
*Discoaster woodringii* Bramlette & Riedel, 1954  
*Geminolithella rotula* (Kamptner, 1956) Backman, 1980  
*Helicosphaera ampliapertura* Bramlette & Wilcoxon, 1967  
*Helicosphaera carteri* (Wallich, 1877) Kamptner, 1954  
*Helicosphaera euphratis* Haq, 1966  
*Helicosphaera mediterranea* Müller, 1981  
*Helicosphaera orientalis* Black, 1971  
*Helicosphaera pacifica* Müller & Brönnimann, 1974  
*Helicosphaera perch-nielseniae* Haq, 1971  
*Helicosphaera scissura* Miller, 1981

- Helicosphaera stalis* Theodoridis, 1984  
*Helicosphaera vedderi* Bukry, 1981  
*Helicosphaera walbersdorfensis* (Müller, 1978) Theodoridis, 1984  
*Helicosphaera waltrans* Theodoridis, 1984  
*Lithostromation perdurum* Deflandre, 1942  
*Minylitha convallis* Bukry, 1973  
*Pyrocyclus inversus* Hay & Towe, 1962  
*Pyrocyclus orangensis* (Bukry, 1971) Backman, 1980  
*Reticulofenestra minuta* Roth, 1970  
*Reticulofenestra minutula* (Gartner, 1967) Haq & Berggren, 1978  
*Reticulofenestra pseudumbilicus* (Gartner, 1967) Gartner, 1969  
 small *Reticulofenestra pseudumbilicus* (5-7 micron)  
 "large" *Reticulofenestra pseudumbilicus* (>12 micron)  
*Scapholithus fossilis* Deflandre in Deflandre & Fert, 1954  
*Sphenolithus abies* Deflandre in Deflandre & Fert, 1954  
*Sphenolithus abies/neoabies* (sensu Rio et al., 1990b)  
*Sphenolithus belemnos* Bramlette & Wilcoxon, 1967  
*Sphenolithus dissimilis* Bukry & Percival, 1971  
*Sphenolithus conicus* Bukry, 1971  
*Sphenolithus heteromorphus* Deflandre, 1953  
*Sphenolithus moriformis* (Bronnimann & Stradner, 1960) Bramlette & Wilcoxon, 1967  
*Sphenolithus neoabies* Bukry & Bramlette, 1969  
*Tetralithoides symeonidesii* (Theodoridis, 1984)  
*Triquetrorhabdulus milowii* Bukry, 1971  
*Triquetrorhabdulus rugosus* Bramlette & Wilcoxon, 1967  
*Triquetrorhabdulus serratus* (Bramlette & Wilcoxon, 1967) Olafsson, 1989  
*Zygrrhabdulus bijugatus* (Deflandre in Deflandre & Fert, 1954) Deflandre, 1959

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