

COMPUTATIONAL METHODOLOGY TO ANALYZE THE EFFECT OF MASS TRANSFER RATE ON ATTENUATION OF LEAKED CARBON DIOXIDE IN SHALLOW AQUIFERS

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ABSTRACT. Exsolution and re-dissolution of CO₂ gas within heterogeneous porous media are investigated using experimental data and mathematical modeling. In a set of bench-scale experiments, water saturated with CO₂ under a given pressure is injected into a 2-D water-saturated porous media system, causing CO₂ gas to exsolve and migrate upwards. A layer of fine sand mimicking a heterogeneity within a shallow aquifer is present in the tank to study accumulation and trapping of exsolved CO₂. Then, clean water is injected into the system and the accumulated CO₂ dissolves back into the flowing water. Simulated exsolution and dissolution mass transfer processes are studied using both near-equilibrium and kinetic approaches and compared to experimental data under conditions that do and do not include lateral background water flow. The mathematical model is based on the mixed hybrid finite element method that allows for accurate simulation of both advection- and diffusion-dominated processes.

KEYWORDS: Compositional flow, two-phase flow, kinetic mass transfer, gas exsolution, gas dissolution.

1. INTRODUCTION

Geologic carbon sequestration has the potential to significantly reduce greenhouse gas emissions [1], but also poses risks to groundwater resources including mobilization of contaminants in shallow aquifers due to leakage of CO₂ from deep storage formations [2]. The extent and severity of the risks depend on complex multiphase flow and transport phenomena that govern the migration of CO₂ through the shallow subsurface. A persistent issue with predicting these processes is the general difficulty of understanding CO₂ phase change (a.k.a. inter-phase mass transfer) within macroscopic porous media systems, which is important in the case of CO₂ due to its high solubility and potential mobility in the gas phase.

If a leakage pathway is encountered, stored CO₂ that is originally supercritical in a deep geologic storage formation may migrate upward due to buoyancy, dissolve into water, exsolve to form a separate gas phase, and eventually re-dissolve into clean water. These interrelated processes are collectively referred to as multiphase evolution, and have recently been studied in various continuum-scale systems. Experimental investigations have identified the factors that control multiphase evolution during one-dimensional (1D) ver-

tical flow [3–5] and have qualitatively addressed the various transport phenomena during two-dimensional (2D) flow [6]. Numerical modeling has been used to show the effects of 1D flow rate [7] and quasi-2D flow paths [8] on multiphase CO₂ evolution processes. However, numerical models have not yet been able to fully explain all of the observations from the experimental studies, particularly of those that occur during 2D flow under the influence of background water flow.

The purpose of this present study is to develop a numerical modeling methodology to assess the factors that control inter-phase CO₂ mass transfer during migration through heterogeneous 2D porous media systems in the shallow subsurface where CO₂ can only exist in the gaseous and dissolved phases. Specifically, we seek to test whether the local equilibrium assumption for mass transfer adequately explains multiphase evolution, and to investigate the effect of lateral water flow on heterogeneity-enhanced gas phase accumulation.

Our approach is to expand the numerical model developed in [9] to include kinetic mass transfer, and then to compare the results of the simulations performed by the model to the laboratory data from experiments that build upon methods developed in

[10]. This methodology forms the basis for another article [11] that investigates multiphase CO₂ plume dynamics in relatively large-scale synthetic aquifer systems. These investigations improve our understanding of the underlying processes that together lead to attenuation of CO₂ within groundwater systems [6].

The main objective of this work is to study the exsolution and dissolution of CO₂ on a small scale using both experiments and numerical modeling. Motivated by fundamental differences in mass transfer in various scenarios studied in [11], our aim is to determine whether a kinetic or a simplified equilibrium mass transfer model is needed to reproduce the results observed in the experiments and how the results are affected by changes in the flow field enforced in the experiments.

2. EXPERIMENT

The experimental methodology is detailed by [10], so it will be only briefly described here. To assess the effects of background (lateral) water flow on multiphase CO₂ migration, two different bench-scale experiments were performed. In each experiment, CO₂-saturated water was injected into the bottom of a tank filled with water-saturated porous media until gas phase formed and accumulated to some steady value. Then, clean water was injected until all of the CO₂ dissolved and migrated out of the tank.

The test system was initially reported by [10], but this study expands upon that methodology by incorporating different experimental conditions. Most importantly, one experiment was performed with water flowing from right to left across the tank. As is shown in Figure 1, a block of low-permeability sand was incorporated into the middle of the tank above the injection ports, and saturation sensors were installed below this block of fine sand where gas phase CO₂ was expected to accumulate after exsolving from the injected CO₂-saturated water. Material properties of the three porous media used are given in Table 3. The saturation pressure, as defined by [3], was 10 kPa for the first experiment and 15 kPa for the second, which led to dissolved CO₂ concentrations that were sufficiently high to cause exsolution in the porous media.

In the first experiment, the constant head devices connected to the regions of Granusil #8 (which were installed to distribute the head evenly across the vertical boundaries of the main sand pack) were positioned at equal elevations. This led to negligible background water flow across the tank, and this case is therefore referred to as the static case. In the second experiment, the left hand constant head device was positioned at an elevation below that of the right-hand one, thus establishing background flow. In the second case, water inflow was supplied to the right-hand constant head device via a peristaltic pump, while the outflow from the left-hand constant head device was routed

into a container placed on a computer-interfaced electronic scale. The scale and the saturation sensors were configured to automatically take measurements at 1-minute intervals.

The resulting data from the experiments were compared against a numerical model of two-phase flow in porous media, which extends upon that of [9]. A major addition to the model is its capability to account for non-equilibrium CO₂ mass transfer between aqueous and gas phases.

3. MATHEMATICAL MODEL

In this section, the mathematical model that describes the two-phase compositional flow in porous media and incorporates the phenomena studied in this work is summarized.

3.1. TWO-PHASE FLOW IN POROUS MEDIA

The governing equations for the two-phase flow in porous media are based on [12–14]. The quantities corresponding to the liquid (wetting) and gas (non-wetting) phases are denoted by indices ℓ and g , respectively.

The mass balance equations for the incompressible liquid and gas phases are given by

$$\rho_\ell \phi \frac{\partial S_\ell}{\partial t} + \rho_\ell \nabla \cdot \vec{v}_\ell = f_\ell, \quad (1a)$$

and

$$\rho_g \phi \frac{\partial S_g}{\partial t} + \rho_g \nabla \cdot \vec{v}_g = f_g, \quad (1b)$$

respectively, where ϕ [–] is the material porosity and S_α [–], ρ_α [kg m⁻³], \vec{v}_α [m s⁻¹], f_α [kg m⁻³ s⁻¹] are the α -phase saturation, density, velocity, and the sink or source term. The velocity \vec{v}_α is given by Darcy's law

$$\vec{v}_\alpha = -\lambda_\alpha K (\nabla p_\alpha - \rho_\alpha \vec{g}), \quad (2)$$

where \vec{g} [m s⁻²] is the gravitational acceleration vector, K [m²] is the intrinsic permeability, p_α [Pa] is the α -phase pressure, $\lambda_\alpha = k_{r\alpha}/\mu_\alpha$ [Pa⁻¹ s⁻¹] denotes the mobility of phase α , where μ_α [Pa s] is the dynamic viscosity, and $k_{r\alpha}(S_\alpha)$ [–] denotes the relative permeability.

The difference between the wetting and non-wetting phase pressures is defined as the capillary pressure $p_c = p_g - p_\ell$ and the Brooks and Corey model [15] is used in the form

$$p_c(S_\ell) = p_d (S_\ell^e)^{-\frac{1}{\lambda}}, \quad (3)$$

where p_d [Pa] is the entry pressure, λ [–] is the pore size distribution index, and S_α^e [–] denotes the effective saturation defined by

$$S_\alpha^e = \frac{S_\alpha - S_{r,\alpha}}{1 - S_{r,g} - S_{r,\ell}}, \quad (4)$$

where $S_{r,\alpha}$ [–] is the residual saturation of phase α , $\alpha \in \{\ell, g\}$.

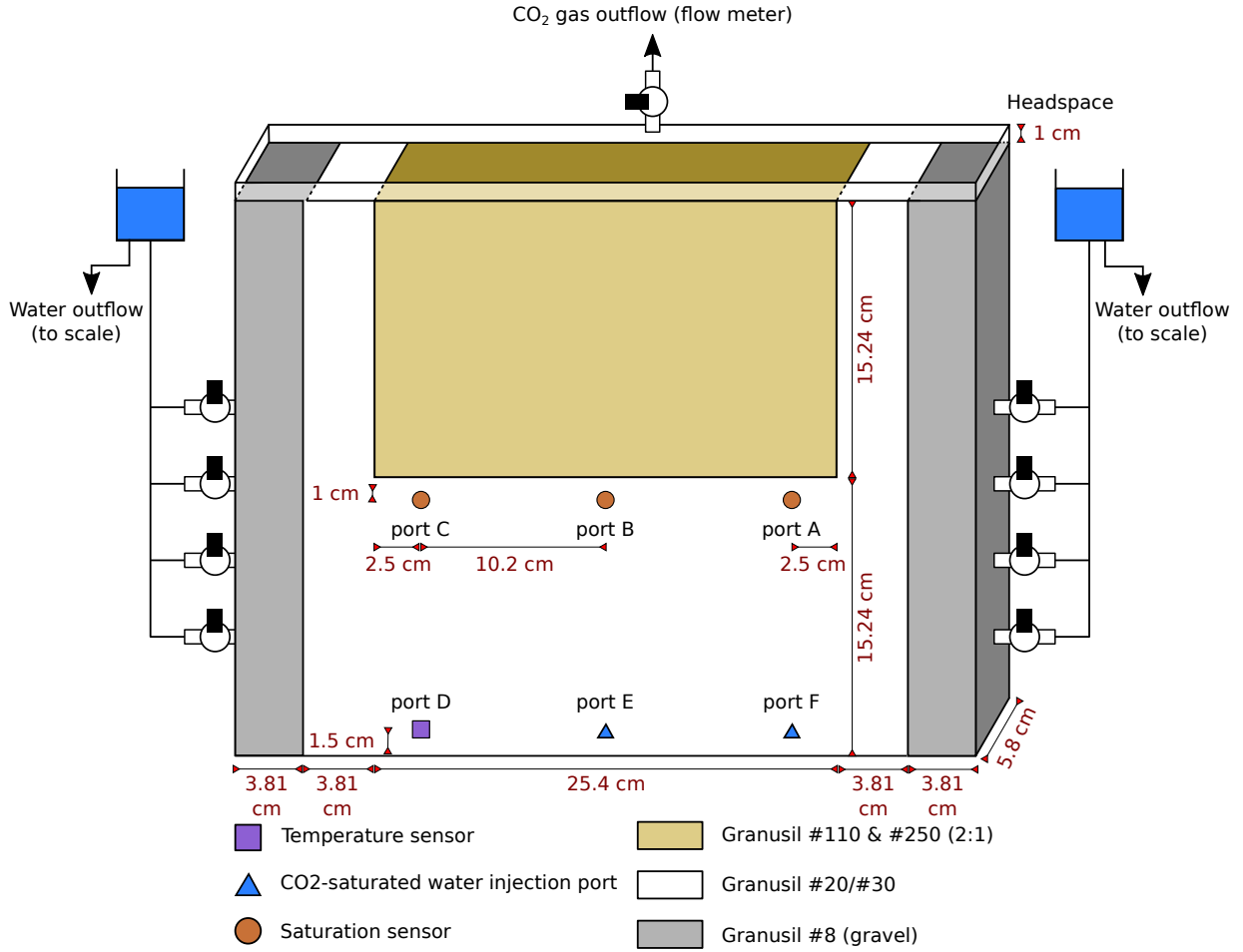


FIGURE 1. Setup for the small tank experiments, adapted from [10].

The Brooks and Corey model parameters are also used in the Burdine's model for the relative permeability functions $k_{r,\ell}$ and $k_{r,g}$, [16], in the form

$$k_{r,\ell}(S_\ell) = (S_\ell^e)^{\frac{2+3\lambda}{\lambda}}, \quad (5)$$

$$k_{r,g}(S_g) = (S_g^e)^2 \left(1 - (1 - S_g^e)^{\frac{2+\lambda}{\lambda}} \right). \quad (6)$$

Based on [7, 11, 17], however, instead of $k_{r,g}$, a modified formula for the gas phase relative permeability function is used in Eq. (2) in the form

$$\tilde{k}_{r,g}(S_g) = \begin{cases} 0, & \text{if } S_g < S_c, \\ k_{r,g}\left(\frac{S_g - S_c}{1 - S_c}\right), & \text{otherwise,} \end{cases} \quad (7)$$

where S_c [-] denotes the critical gas saturation.

3.2. COMPONENT TRANSPORT

The gas phase is considered as a single component (pure CO₂), whereas the liquid phase is assumed to be a two component mixture (water and CO₂). The compositional balance equation for CO₂ dissolved in

the liquid phase is included into (1) and (2) as

$$\phi \rho_\ell \frac{\partial(S_\ell X)}{\partial t} + \rho_\ell \nabla \cdot (X \vec{v}_\ell - \tau_\ell \phi S_\ell D_\ell \nabla X) = f_X, \quad (8)$$

[14], where X [-] is the mass fraction of CO₂, f_X [kg m⁻³ s⁻¹] is the sink or source term, D_ℓ [m² s⁻¹] is the free molecular diffusion of CO₂ in water, $D_\ell = 1.92 \cdot 10^{-9}$ m² s⁻¹, and τ_ℓ [-] is the tortuosity given by $\tau_\ell = \phi^{1/3} S_\ell^{7/3}$ based on [18].

3.3. KINETIC MASS TRANSFER

Based on [19], the kinetic mass transfer of CO₂ between both phases (i.e., the dissolution and exsolution processes) is represented by

$$-f_g = f_\ell = f_X = k(C_s - X \rho_\ell), \quad (9)$$

where k [s⁻¹] is the lumped mass transfer rate coefficient and C_s [kg m⁻³] is the saturated CO₂ concentration in water at the relevant pressure and temperature given by Henry's law in the form

$$C_s = \frac{p_g}{K_H} M_g, \quad (10)$$

where M_g [kg mol⁻¹] is the molar mass of CO₂, $M_g = 44.01$ g mol⁻¹, and K_H [Pa mol⁻¹ m³] is Henry's constant for which the Van't Hoff equation is employed in the form

$$K_H = K_{H,ref} e^{-C \left(\frac{1}{T} - \frac{1}{T_{ref}} \right)} \quad (11)$$

where T [K] is the temperature, $K_{H,ref}$ is the value of Henry's constant at a reference temperature T_{ref} [K], and C [K] is the gas-specific constant, i.e., $K_{H,ref} = 2979.97$ [Pa mol⁻¹ m³], $T_{ref} = 298.15$ K, and $C = 2400$ K [20]. The average temperatures in the static case and background flow experiments were 37°C and 26°C, respectively.

4. NUMERICAL METHOD AND IMPLEMENTATION

The governing equations are solved using a general numerical solver based on the mixed-hybrid finite element method described in [9]. The mixed-hybrid finite element method combines velocity discretizations in the lowest order Raviart–Thomas–Nédélec space with piecewise constant approximations for the scalar variables. The numerical method can be used for accurate simulation of degenerate diffusion or advection-dominated problems like the one discussed here.

4.1. NUMERICAL METHOD

The numerical method solves a system of n partial differential equations in the coefficient form in a d -dimensional polygonal domain $\Omega \subset \mathbb{R}^d$ and a time interval $[0, t_{fin}]$:

$$\sum_{j=1}^n N_{i,j} \frac{\partial Z_j}{\partial t} + \sum_{j=1}^n \tilde{u}_{i,j} \cdot \nabla Z_j + \nabla \cdot (m_i(\vec{q}_i + \vec{w}_i)) = f_i, \quad (12a)$$

with

$$\vec{q}_i = - \sum_{j=1}^n D_{i,j} \nabla Z_j, \quad (12b)$$

where $Z_j = Z_j(t, \vec{x})$, $j = 1, 2, \dots, n$, represent the unknown variables, $\vec{x} \in \Omega$, $t \in [0, t_{fin}]$. Eq. (12) is further supplemented with Dirichlet or Neumann boundary conditions, or its combination on different parts of the boundary [9].

Here we restrict ourselves only the case $d = 2$ which is relevant for the problems presented in the paper. In brevity, the mixed hybrid finite element discretization approximates the unknown scalar functions Z_j in the space of piecewise constant functions, i.e.,

$$Z_j \approx \sum_{A \in \mathcal{A}_h} Z_{j,A} \varphi_A \quad (13)$$

and the vector function \vec{q}_i from Eq. (12b) in the lowest order Raviart–Thomas–Nédélec space RTN₀ as

$$\vec{q}_i \approx \sum_{A \in \mathcal{A}_h} \chi_A \sum_{E \in \mathcal{E}_A} q_{i,A,E} \vec{\omega}_{A,E}, \quad (14)$$

where \mathcal{A}_h denotes the set of triangles discretizing the computational domain Ω , h is the mesh size defined as the largest ball diameter circumscribed around elements in \mathcal{A}_h , \mathcal{E}_A is the set of all sides of an element $A \in \mathcal{A}_h$, φ_A are the piecewise constant basis functions defined by $\varphi_A = \chi_A$, where χ_A is the characteristic function of A defined by

$$\chi_A(\vec{x}) = \begin{cases} 1 & \vec{x} \in A, \\ 0 & \vec{x} \notin A, \end{cases} \quad (15)$$

and $\vec{\omega}_{A,E}$ are the vector basis functions of RTN₀(A), [9].

Eq. (12b) and the discretization given by Eq. (14) allow to express the coefficients $q_{i,A,E}$ as

$$q_{i,A,E} = \sum_{j \in \sigma_{i,A}} \left(b_{i,j,A,E} Z_{j,A} - \sum_{F \in \mathcal{E}_A} b_{i,A,E,F} Z_{j,F} \right), \quad (16)$$

where $Z_{j,E}$ are traces of Z_j on side $E \in \mathcal{E}_A$, $\sigma_{i,A} \subseteq \{1, 2, \dots, n\}$ denotes the set of all indices j for which $D_{i,j}$ is non-zero on element A and the definition of coefficients $b_{i,j,A,E}$ and $b_{i,A,E,F}$ are given in [9].

The approximation of \vec{q}_i in RTN₀ given by Eqs. (14) and (16) together with the piecewise constant approximation of Eq. (12a) allow to express the cell-averages $Z_{j,A}$ as a local linear combination of $Z_{j,E}$, $E \in \mathcal{E}_A$, on each element $A \in \mathcal{A}_h$, see [9] for details. Hence, $q_{i,A,E}$ can be expressed solely as linear function of traces $Z_{j,E}$, $\forall E \in \mathcal{E}_A$.

Across all interior sides E of the triangulation \mathcal{A}_h , the balance of conservative fluxes is given by

$$\sum_{\{A:E \in \mathcal{E}_A\}} m_{i,E} (\vec{q}_{i,A,E} + \vec{w}_{i,A,E}) = 0, \quad (17)$$

where a single (upwinded) value $m_{i,E}$ is assumed at E , see [9], and, thus, it can be eliminated from Eq. (17) if it has a non-zero value to produce

$$\sum_{\{A:E \in \mathcal{E}_A\}} \vec{q}_{i,A,E} + \vec{w}_{i,A,E} = 0. \quad (18)$$

When $m_{i,E} = 0$ in Eq. (17) for some side E , Eq. (18) is also considered to assure that the resulting system of linear equations, given in the vector form as

$$\mathbf{M}^k \vec{Z}^{\ell+1} = \vec{b}^\ell, \quad (19)$$

is non-singular. In Eq. (19), \mathbf{M} is a sparse, positive definite matrix [9] and \vec{b} is the right-hand-side, both evaluated on the previous time level ℓ , and \vec{Z} is the vector containing side-traces $Z_{j,E}$ on the next time level $\ell + 1$ for all $j = 1, 2, \dots, n$ and all sides E in \mathcal{A}_h .

4.2. PROBLEM FORMULATION

The system of governing equations given by Eqs. (1), (2), (8), and (9) is represented by (12) using $n = 3$,

Parameter	Units	liquid	gas
		H ₂ O	CO ₂
Density	ρ [kg m ⁻³]	997.78	1.98
Dyn. viscosity	μ [10 ⁻⁵ Pa s ⁻¹]	87.2	1.48

TABLE 1. Properties of fluids.

	Units	Mesh	Mesh	Mesh
		Δ_1	Δ_2	Δ_3
Elements		5 494	18 587	72 949
Mesh size h	[mm]	5.55	2.78	1.43

TABLE 2. Properties of meshes.

$d = 2$, $Z_1 = p_c$, $Z_2 = p_g$, $Z_3 = X$, and

$$(N_{i,j})_{i,j \in \hat{3}} = \begin{pmatrix} -\phi \rho_\ell \frac{dS_\ell}{dp_c} & 0 & 0 \\ -\phi \rho_g \frac{dS_\ell}{dp_c} & 0 & 0 \\ 0 & 0 & \phi S_\ell \rho_\ell \end{pmatrix}, \quad (20)$$

$$(\vec{u}_{i,j})_{i,j \in \hat{3}} = \begin{pmatrix} \vec{0} & \vec{0} & \vec{0} \\ \vec{0} & \vec{0} & \vec{0} \\ \vec{0} & \vec{0} & \rho_\ell \vec{v}_\ell \end{pmatrix}, \quad (21)$$

$$(m_i)_{i \in \hat{3}} = \begin{pmatrix} \rho_\ell \frac{\lambda_\ell}{\lambda_\ell + \lambda_g} \\ \rho_g \frac{\lambda_g}{\lambda_\ell + \lambda_g} \\ \rho_\ell \end{pmatrix}, \quad (22)$$

$$(D_{i,j})_{i,j \in \hat{3}} = \begin{pmatrix} (\lambda_\ell + \lambda_g)K & -\lambda_\ell + \lambda_g K & 0 \\ 0 & (\lambda_\ell + \lambda_g)K & 0 \\ 0 & 0 & \tau \phi S_\ell D_\ell \end{pmatrix}, \quad (23)$$

$$(\vec{w}_i)_{i \in \hat{3}} = \begin{pmatrix} -(\lambda_\ell + \lambda_g) \rho_\ell K \vec{g} \\ \lambda_\ell + \lambda_g \rho_g K \vec{g} \\ \vec{0} \end{pmatrix}, \quad (24)$$

$$(f_i)_{i \in \hat{3}} = \begin{pmatrix} -f_\ell \\ f_g \\ f_\ell - X f_\ell \end{pmatrix}, \quad (25)$$

where $\hat{3} = \{1, 2, 3\}$.

The computational domain Ω depicted in Figure 2 is discretized using three gradually refined triangular meshes generated by `gmsh` [21]. The mesh properties are listed in Table 2 and in Figure 2, the coarsest mesh is shown.

The numerical method is implemented in an in-house computer code using `C++`. The applicability of the numerical method for heterogeneous porous media is further discussed in [22] together with parallel implementations of the method on GPU [9] or on CPU using MPI [23].

5. RESULTS AND DISCUSSION

5.1. COMPUTATIONAL STUDY

After the physical domain and boundaries were set up to mimic the experiments, several simulations were performed for each experimental case (i.e., the static

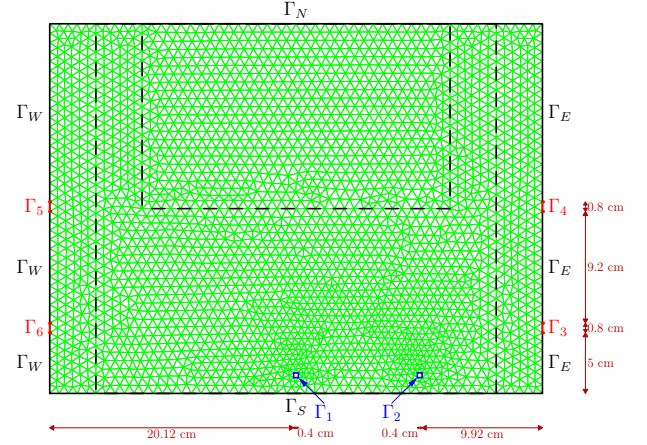


FIGURE 2. Computational domain with coarse mesh based on Figure 1.

one and the one with background flow), using several different values for k . The boundary conditions on the model included hydrostatic pressure in the inflow/outflow ports on the left and right-hand side of the tank (see Figure 1). For the experiment with background flow, the pressure on the right-hand side was increased to enforce a lateral pressure gradient and thus background flow through the tank. The pressure difference across the tank was fitted to match the measured outflow rates from the experiment. Injection of CO₂-saturated and clean water was represented by a Neumann boundary condition at the injection port. No flow boundary conditions were prescribed along the remaining parts of the computational domain boundary.

Several additional simulations were performed to investigate the effect of different background water flow rates on gas accumulation below the heterogeneity. All of these simulations were performed with a high (near-equilibrium) mass transfer rate of $k = 1 \text{ s}^{-1}$. The simulation representing the experiment with background flow was conducted with a pressure difference of 5 Pa, while the other simulations were performed with pressure differences of 2, 4, 6, 7, 8, and 9 Pa.

5.2. NUMERICAL CONVERGENCE

The mesh effects on the numerical solution are investigated using three gradually refined triangular meshes described in Table 2. The gas saturation profiles at Ports A, B, and C for $k = 0.5 \text{ s}^{-1}$ and $S_c = 0.2$ for both static and background flow cases are compared in Figure 3. The amount of gas simulated present in ports A, B, and C is slightly overestimated for coarser meshes with respect to finer meshes, however, these results are expected due to the first order of convergence of the numerical method as reported in [9].

5.3. DISCUSSION

In this section, the comparison between the experimental data and the numerical results is presented

Parameter	Units	Granusil #8	Granusil #20/30	2:1 mixture of Granusil #110 and #250
Porosity	ϕ [-]	0.4	0.32	0.35
Intrinsic permeability	K [m ²]	1×10^{-9}	2.30×10^{-10}	6.36×10^{-14}
Pore size distribution index	λ [-]	4.275	7.33	5.35
Entry pressure	p_d [Pa]	600	1200	8027
Residual liquid phase saturation	$S_{r,\ell}$ [-]	0.084	0.084	0.17
Residual gas phase saturation	$S_{r,g}$ [-]	0	0	0

TABLE 3. Material properties.

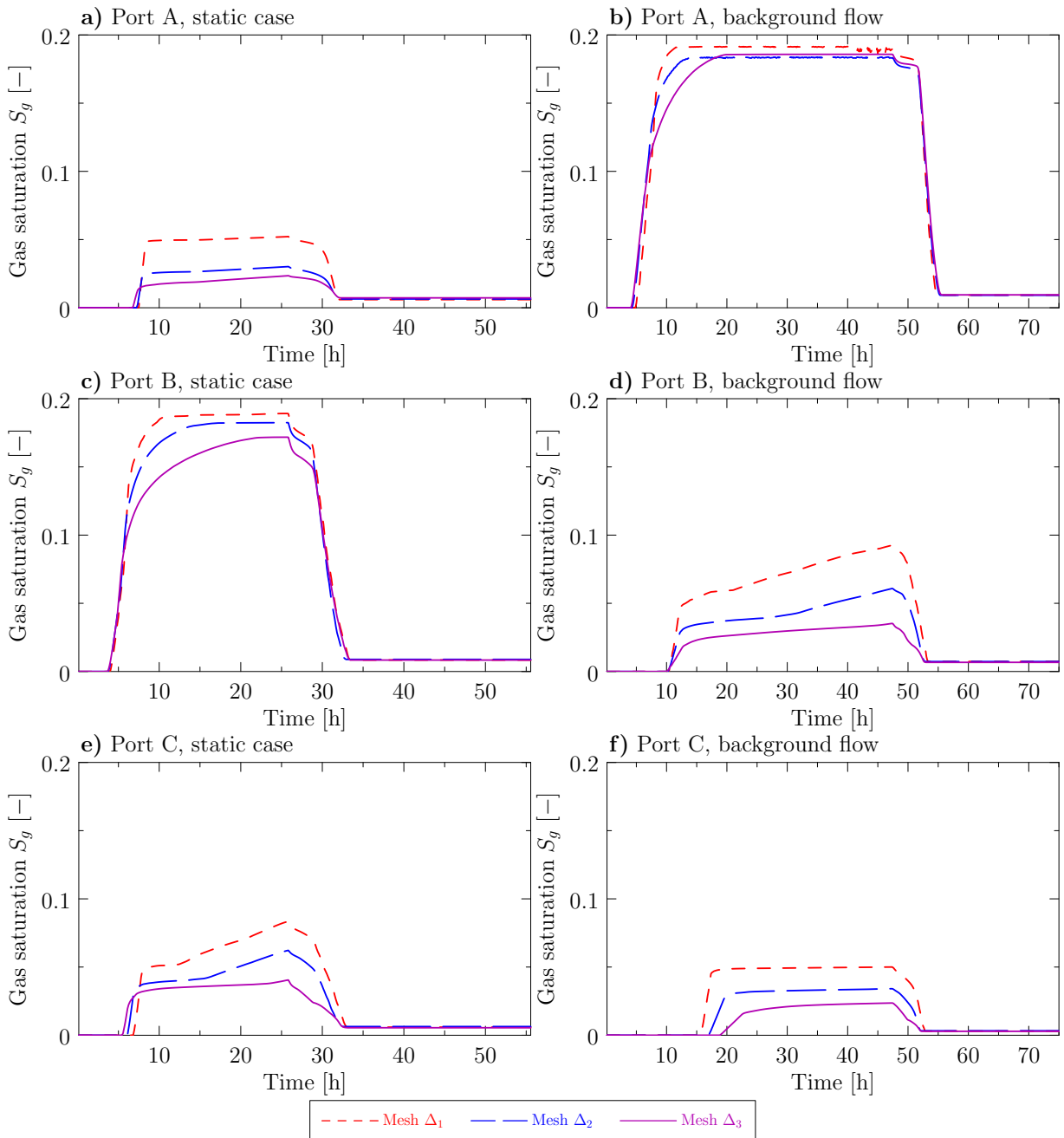


FIGURE 3. Illustration of mesh resolution effects on the numerical solution of the gas saturation evolution for $k = 0.5 \text{ s}^{-1}$ and $S_c = 0.2$ at the sampling ports A (a, b), B (c, d), and C (e, f) for both static (a, c, e) and background flow cases (b, d, f). The mesh properties are given in Table 2.

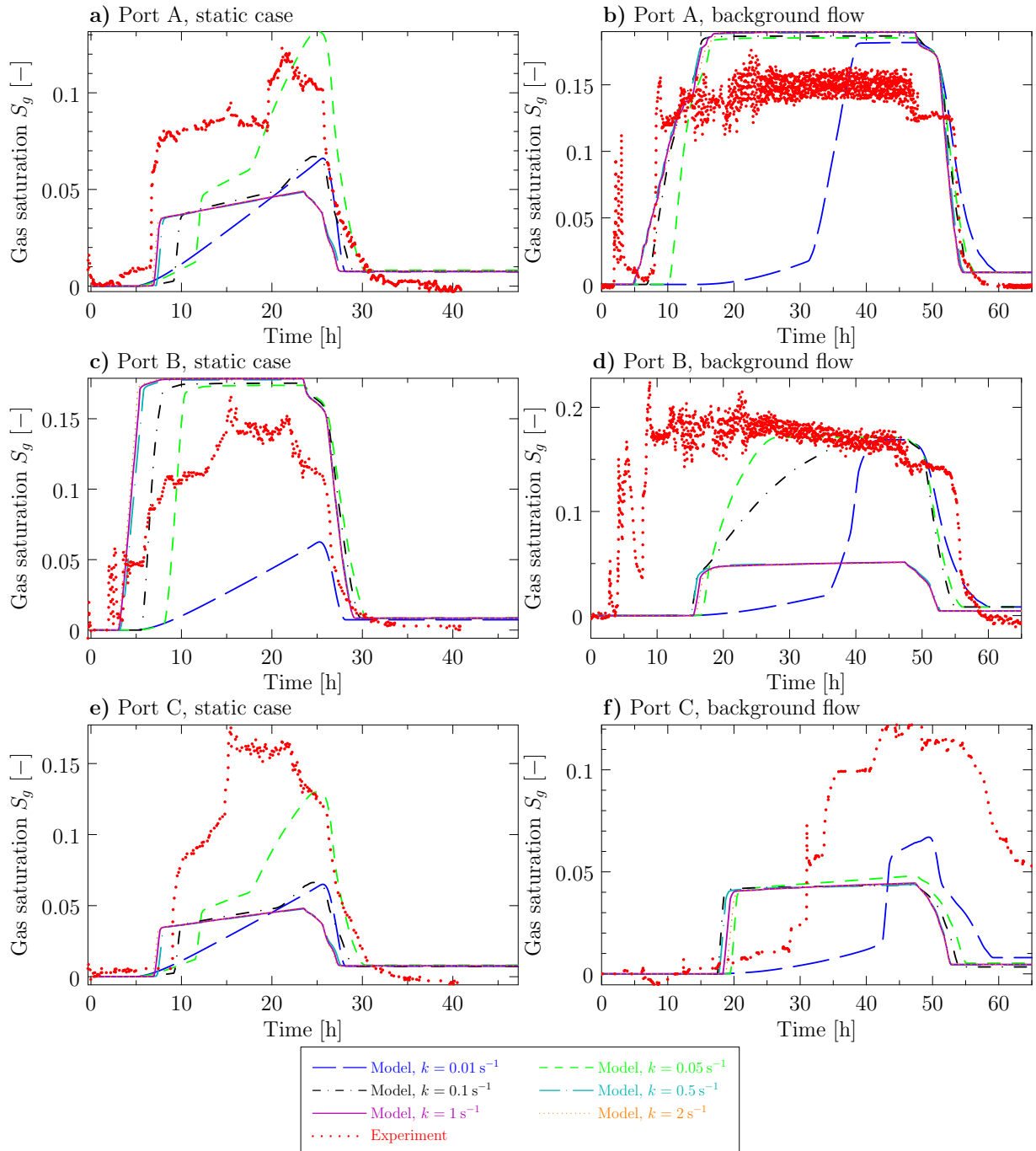


FIGURE 4. Results of the experiments and numerical simulations for (a) the static case and (b) the case with background flow.

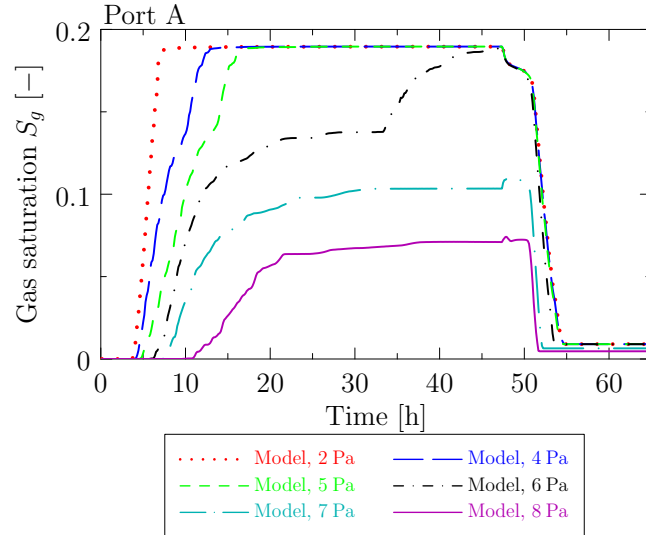


FIGURE 5. Results of the simulations performed with various pressure differences applied across the tank. The experiment with background flow corresponds to the simulation with a pressure difference of 5 Pa.

and the mass transfer and accumulation processes are discussed for both static and background flow cases.

Figure 4 shows the results of gas saturation from the experiments and simulations for the static case and the case with background flow. The experimental data shown in (a) is from port B (directly above the injection port, as CO_2 -saturated water was injected into the left injection port in this case), while that in (b) is from port A (also above the injection port, as CO_2 -saturated water was injected into the right injection port in this case). Both experiments show the accumulation of gas phase CO_2 , then re-dissolution during the injection of clean water after the CO_2 saturations had stabilized.

The simulation results for various choices of k show that the best match with experimental data is achieved for rather large values of the mass transfer coefficient and that the simulated results are almost the same for those greater than a certain threshold value of 0.1 to 0.5 s^{-1} . This indicates that the mass transfer processes are at near-equilibrium and that the background flow is not strong enough to substantially affect the CO_2 exsolution and dissolution rates in these cases. On the other hand, the numerical results show that the presence of the background flow slows down the accumulation dynamics below the fine layer as indicated by different slopes of the saturation curves in Figure 4c compared to Figure 4b. Note that because different injection ports were used in each experiment, the corresponding ports are shifted by one, i.e., results from port B for the static case are compared to the results from port A for the background flow case (see the gas saturation and CO_2 mass fraction spatial evolution shown in Figures 6 and 7).

Further investigation of the background flow effect on the exsolution process can be done by analyzing the results of the simulations performed with different background flow rates, as shown in Figure 5. The re-

sults show a wide range of different gas accumulation behaviors, with higher flow rates leading to lower and slower exsolution due to different gas phase distribution around the injection ports. This indicates that the flow field significantly affects the multiphase CO_2 evolution, however, the experimental data for these scenarios would be needed to quantify the impact.

While the analysis above indicated that near-equilibrium behavior occurred in the experiments, further experimental data are needed to determine generally in which cases the equilibrium mass transfer simplification is valid. Both experiments considered in this work are on a small scale and the measurement ports are close to the injection. Also, the difference between the static and background flow cases is not very substantial. Addressing these remaining knowledge gaps is one of the primary goals in [11], where non-equilibrium mass transfer behavior is observed in a large laboratory-scale system with background flow, and fundamentally different results are observed as compared with predominantly one-dimensional flows. This indicates that, for the decision between the kinetic and equilibrium models, flow field and scale both need to be taken into account. Additionally, the flow field and scale are not the only factors controlling the multiphase evolution, as also indicated in [11], which showed that temperature fluctuations can also significantly affect mass transfer processes.

6. CONCLUSIONS

To assess the potential risks of groundwater contamination via leakage of stored CO_2 from deep geologic carbon sequestration sites, we must be able to predict multiphase CO_2 transport through shallow aquifers. To aid in this effort, an innovative numerical model was applied and tested against data from well-controlled, small-scale laboratory experiments. Beyond conventional continuum-based two-phase flow

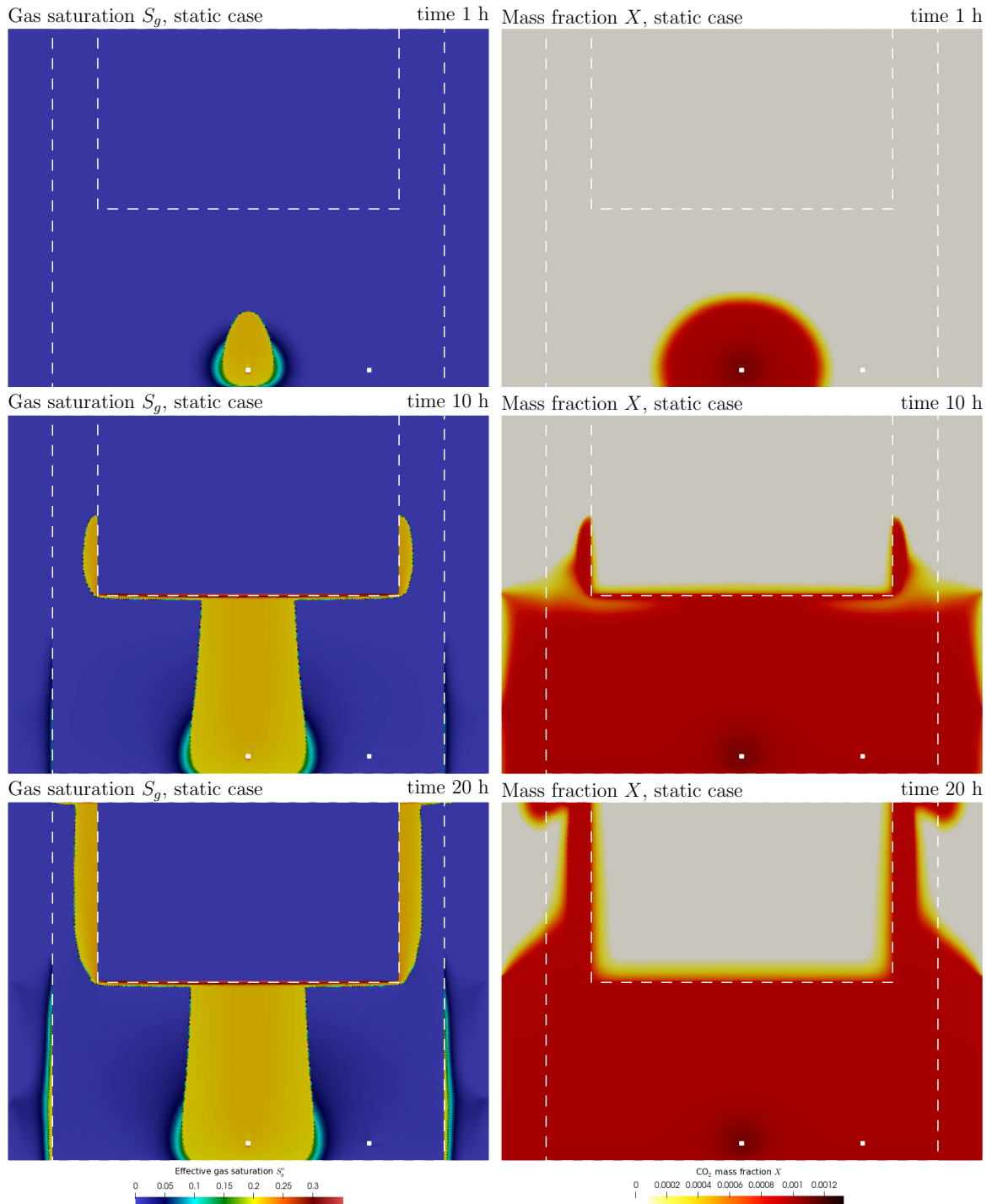


FIGURE 6. Spatial distribution of S_g (left) and X (right) for the static case at 1 h, 10 h, and 20 h (from top to bottom), computed using $k = 0.5 \text{ s}^{-1}$ and $S_c = 0.2$.

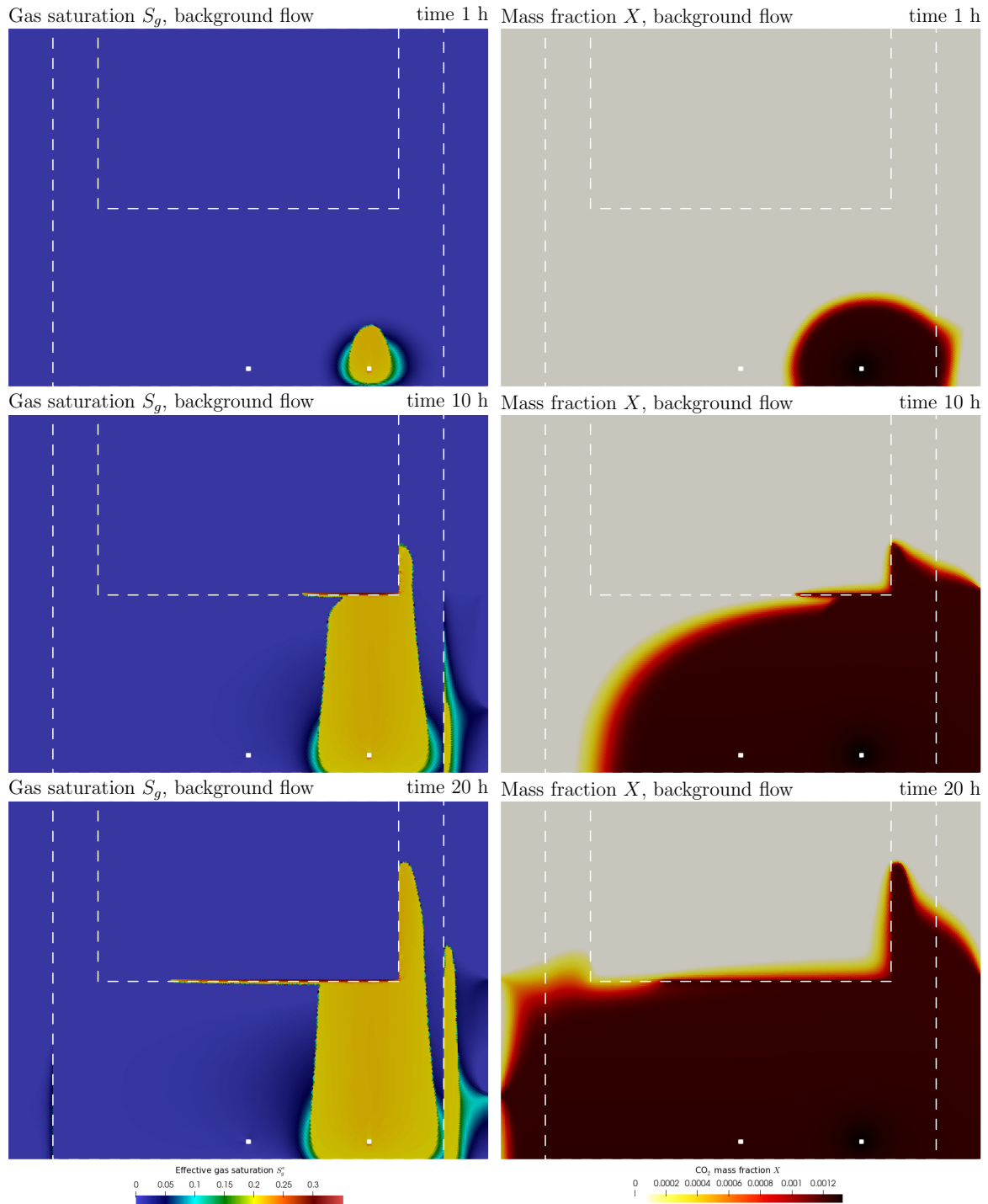


FIGURE 7. Spatial distribution of S_g (left) and X (right) for the background flow case at 1 h, 10 h, and 20 h (from top to bottom), computed using $k = 0.5 \text{ s}^{-1}$ and $S_c = 0.2$.

in porous media, the model incorporated the ability to simulate non-equilibrium CO₂ mass transfer between the aqueous and gaseous phases during exsolution and dissolution processes.

The experimental data used in this work clearly showed equilibrium or near-equilibrium behavior for both steady and background flow cases. These findings are in contrast with the non-equilibrium behavior observed on a larger scale for the scenarios with background flow. This indicates that the mass transfer for the similar scenarios is not controlled by the flow field only. Further investigation of the mass transfer is needed to understand for which scenarios the kinetic model is needed and when the equilibrium assumption is sufficient. The computational methodology developed here could be used to assess both equilibrium and non-equilibrium processes in these future scenarios.

7. ACKNOWLEDGMENTS

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